

CROCO for wave-resolving nearshore dynamics

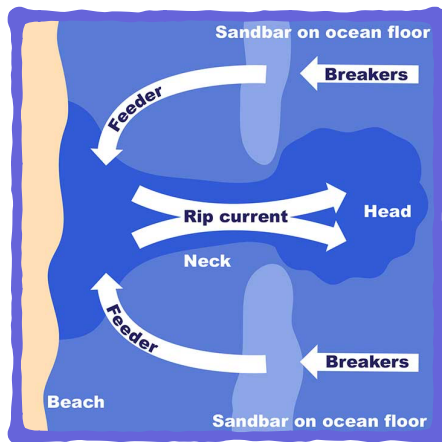
CROCO dev, 2023

P. Marchesiello, S. Treillou, F. Auclair, L. Debreu, J.C. McWilliams,
R. Almar, R. Benshila, F. Dumas

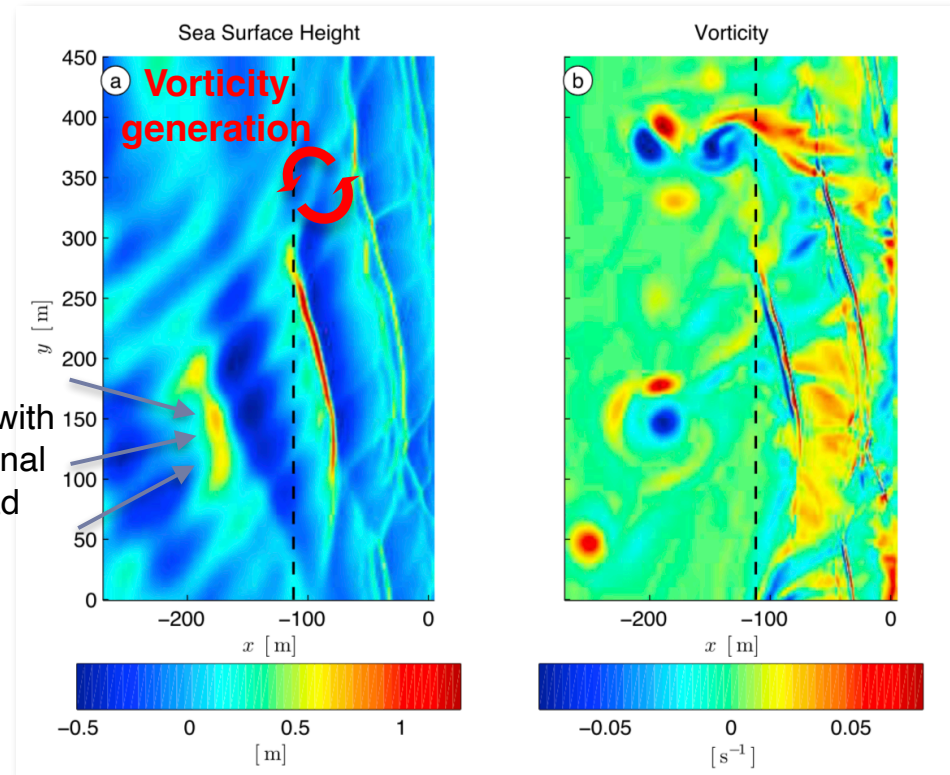
Surfzone eddies are largely due to individual short-crested waves (Peregrine, 1998)



Swimmers caught in a rip current at Haeundae Beach,



Waves with
directional
spread



2D wave-resolving Boussinesq model
(Feddersen et al., 2011)

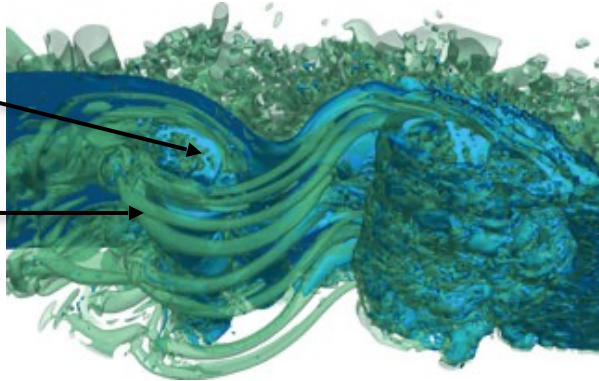
Coastal and estuarine 3D wave-resolving models



- ▶ VOF/SPH LES models: “Direct” turbulence

Roller vortices

Rib vortices

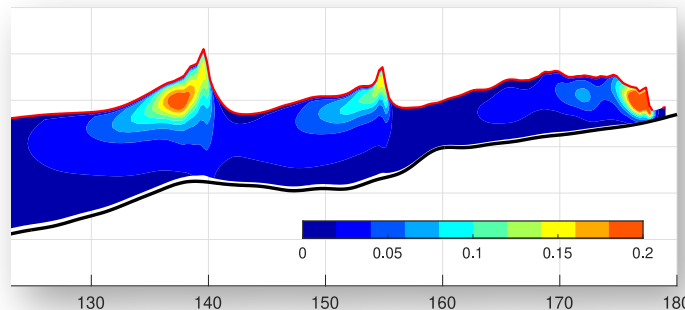


Time scale < wave period

(Lubin & Glockner, 2015)

- ▶ Free-surface RANS models: “Instabilities of the undertow”

CROCO
NHWAVE
SWASH

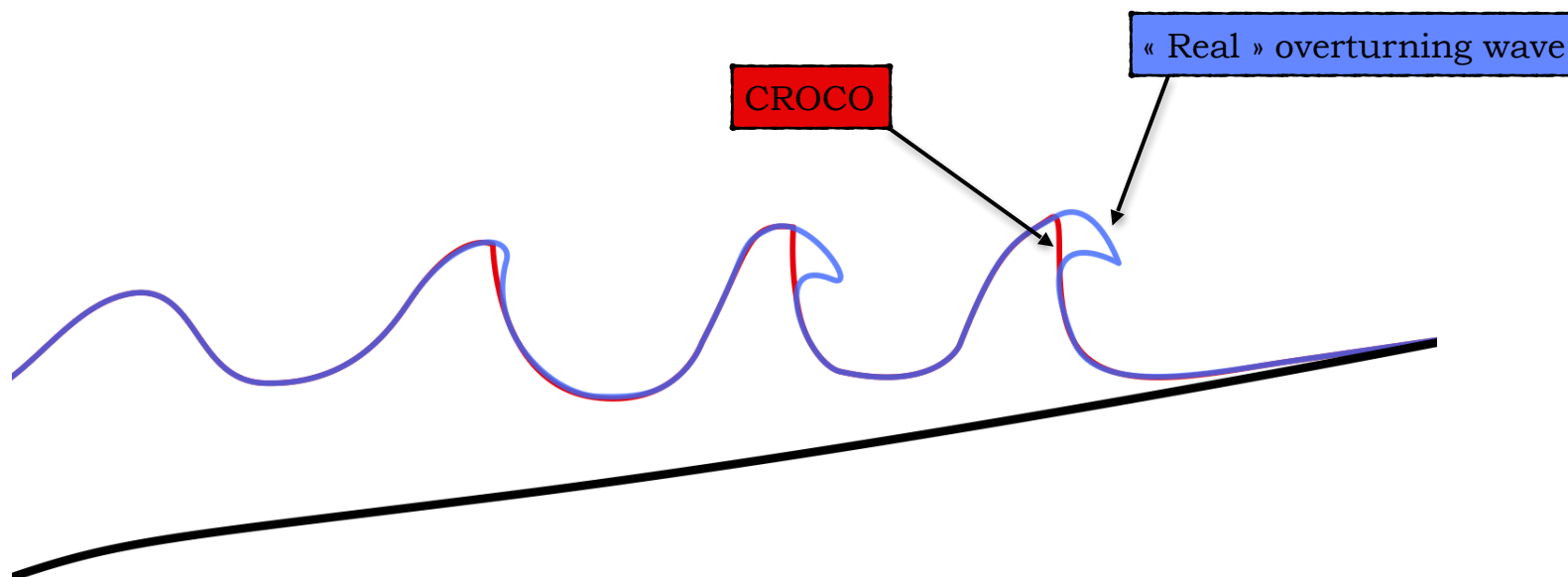


Time scale > wave period

Breaking waves are
treated as shocks

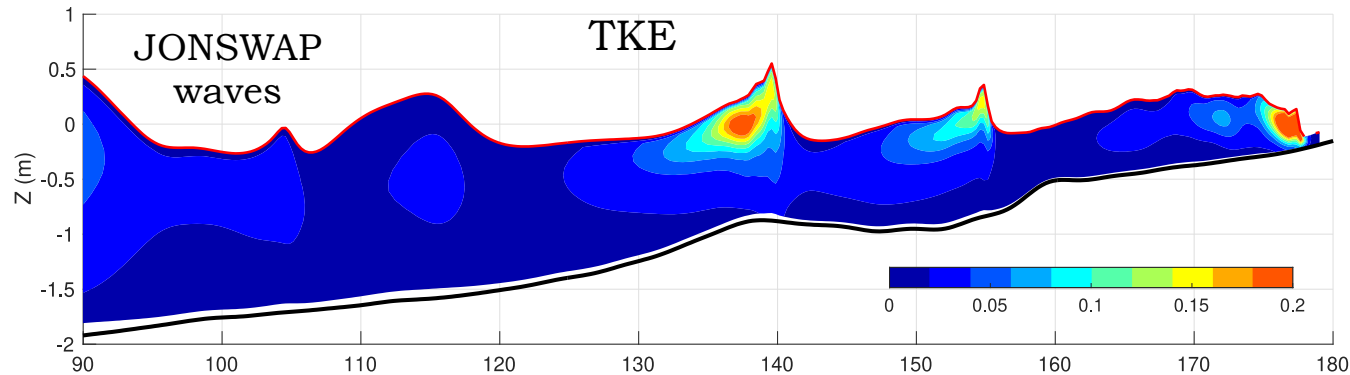
(Marchesiello et al., 2021)

How are breaking waves treated in CROCO



- ✓ Breaking waves are treated as **bores** or shocks with a **shock-capturing scheme** (WENO5)... see Bonneton JFM 2023
- ✓ When the wave is too steep, the model naturally transfers energy to the **mean currents** (about 90%) and to the **sub-grid turbulence** (via turbulence closure).

GLS subgrid scale turbulence



$$\nu_t \sim k^{\frac{1}{2}} l \sim kT \sim \frac{k}{\omega} \sim \frac{k^2}{\epsilon}$$

[m²/s]

k turbulent kinetic energy
l mixing length
T correlation time scale
 ω frequency of turbulence decay (dissipation rate)
 ϵ turbulence dissipation



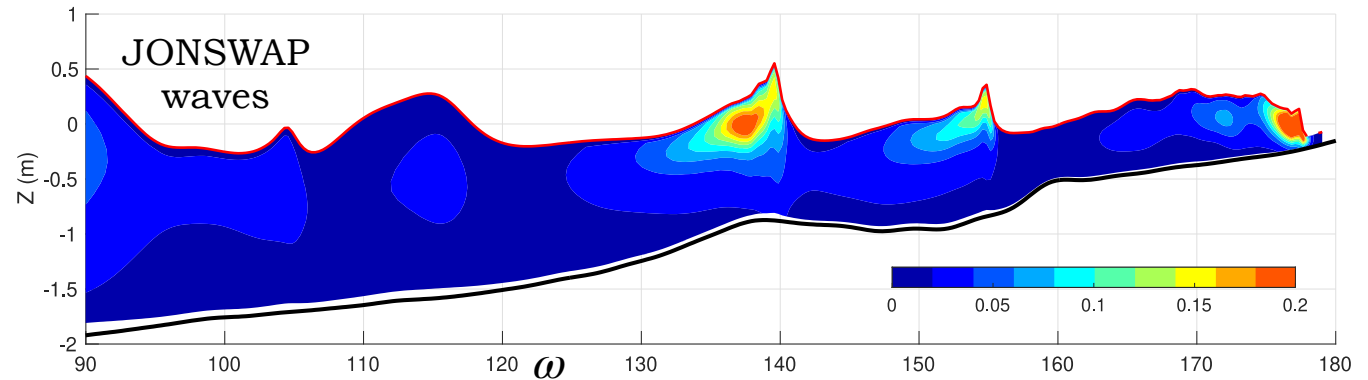
Delta Flume (Deltares)

Validation with flume experiments

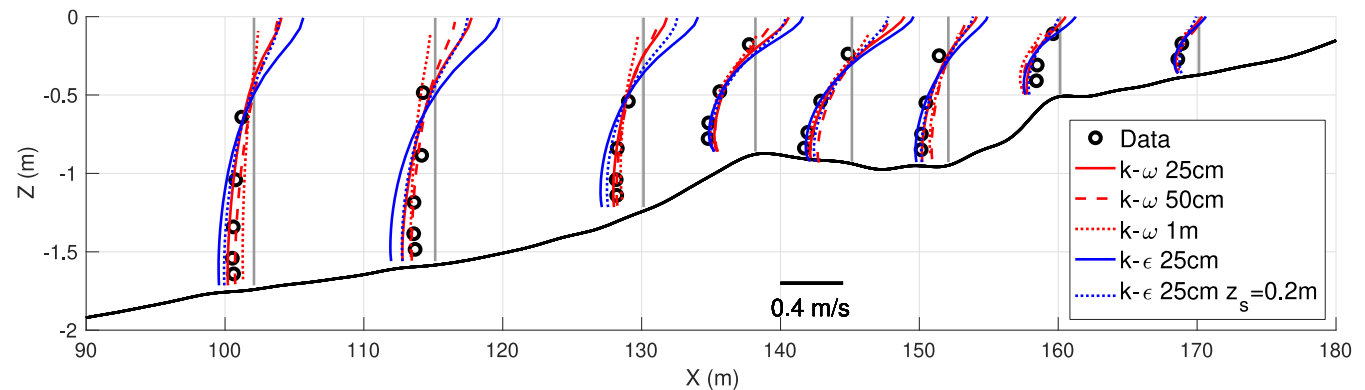
LIP-11D (1B) - Roelvink & Reniers (1995)

► [Marchesiello et al. \(2021, 2022\)](#)

TKE



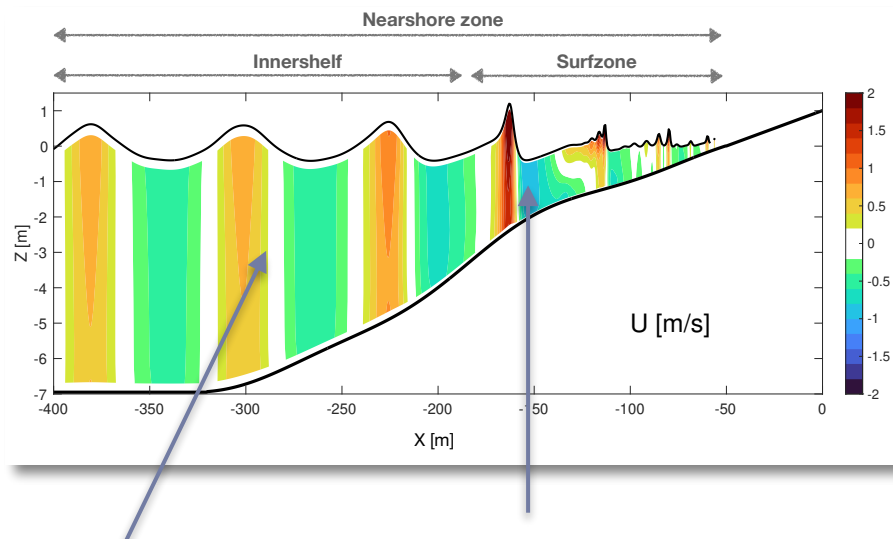
Mean U
Test of resolution
and turbulence
model



Good performances of GLS two-equation turbulence models, especially $k-\omega$

Stabilized GLS turbulence closure

(Marchesiello & Treillou 2023, Larsen & Fuhrman 2018, Mayer & Madsen 2000)



Near-potential flow
(irrotational) region
of non-breaking
waves
 $\frac{\Omega}{S}$ small

Breaking waves

$\frac{\Omega}{S}$ large

Linearized k- ω model

$$\begin{cases} \frac{\partial k}{\partial t} = \overset{\text{Shear production}}{P} - \overset{\text{Dissipation}}{\epsilon} \\ \frac{\partial \omega}{\partial t} = \frac{\omega}{k} (\overset{\text{Dissipation rate } \sim \frac{\epsilon}{k}}{c_{\omega 1} P - c_{\omega 2} \epsilon}) \end{cases}$$

Shear Production $P = 2\nu_t S^2$

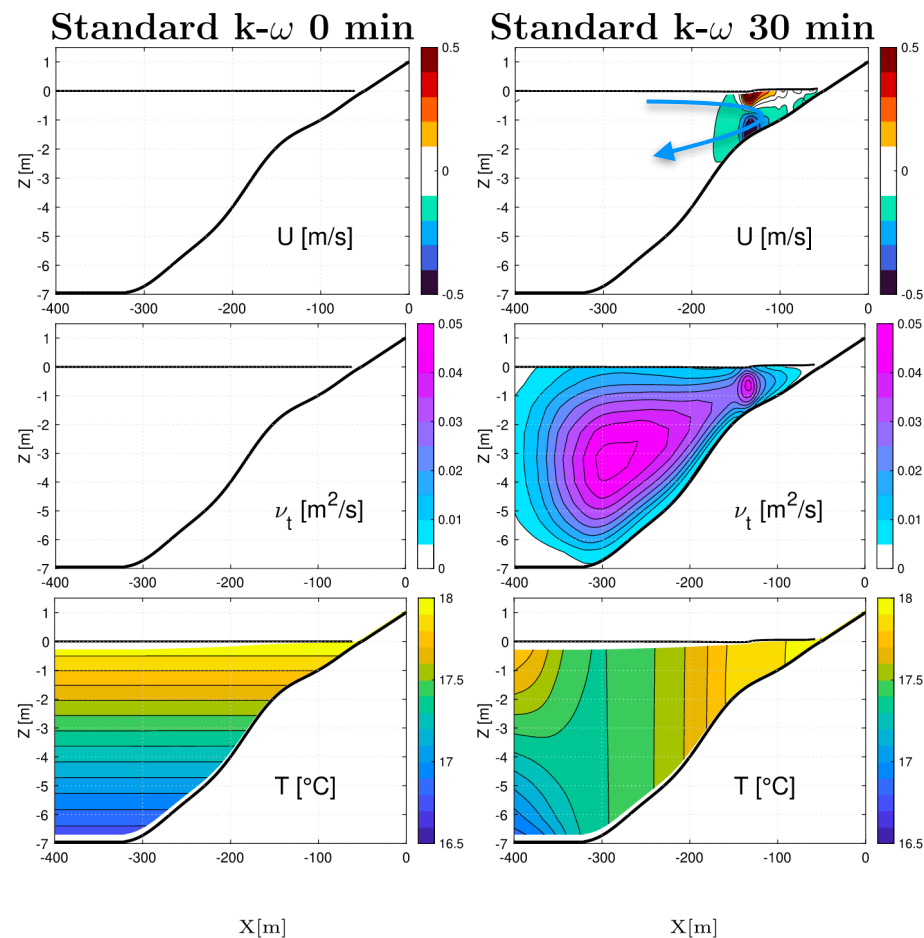
Eddy $\nu_t = \frac{\omega}{k} = c_\mu \frac{k^2}{\epsilon}$ 3D

Strain $S^2 = S_{ij} S_{ij} \quad S_{ij} = \frac{1}{2} \left(\frac{\partial u_i}{\partial x_j} + \frac{\partial u_j}{\partial x_i} \right)$

Rotation rate (vorticity) $\Omega^2 = \Omega_{ij} \Omega_{ij} \quad \Omega_{ij} = \frac{1}{2} \left(\frac{\partial u_i}{\partial x_j} - \frac{\partial u_j}{\partial x_i} \right)$

Stabilized GLS turbulence closure

(Marchesiello & Treillou 2023, Larsen & Fuhrman 2018)



$$\begin{cases} \frac{\partial k}{\partial t} = P - \epsilon \\ \frac{\partial \omega}{\partial t} = \frac{\omega}{k} (c_{\omega 1} P - c_{\omega 2} \epsilon) \end{cases}$$

Shear production (red arrow pointing to P)
Dissipation (red arrow pointing to ϵ)
Dissipation rate (red arrow pointing to $\frac{\omega}{k}$)

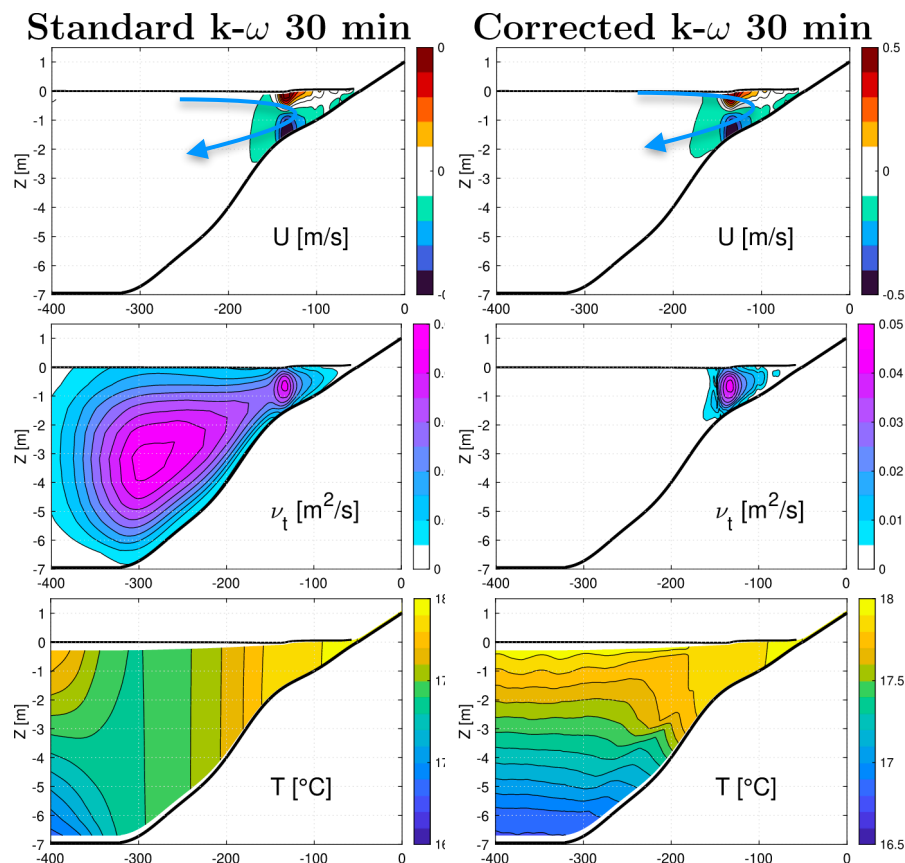
$$P = 2\nu_t S^2 \quad \nu_t = c_\mu \frac{k^2}{\epsilon}$$

Instability of the 2-equation turbulence closure equations :

- ✓ **Overmixing** under non-breaking waves
- ✓ Destruction of innershelf stratification

Stabilized GLS turbulence closure

(Marchesiello & Treillou 2023, Larsen & Fuhrman 2018)



Akin to realizability constraints for stagnation point anomalies (Durbin 2009):

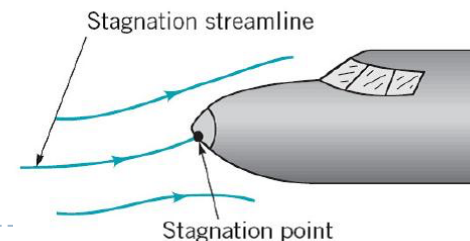
$$\begin{cases} \frac{\partial k}{\partial t} = P - \epsilon \\ \frac{\partial \omega}{\partial t} = \frac{\omega}{k} (c_{\omega 1} P - c_{\omega 2} \epsilon) \end{cases}$$

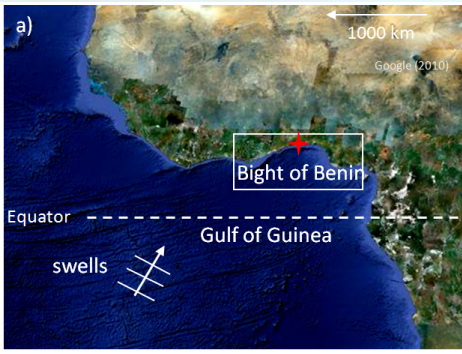
$P = 2\nu_t S^2 \quad \nu_t = c_\mu \frac{k^2}{\epsilon^*}$

$$\epsilon^* = \max \left(1, \lambda \frac{S^4}{\Omega^4} \right) \epsilon$$

same for all
GLS
models

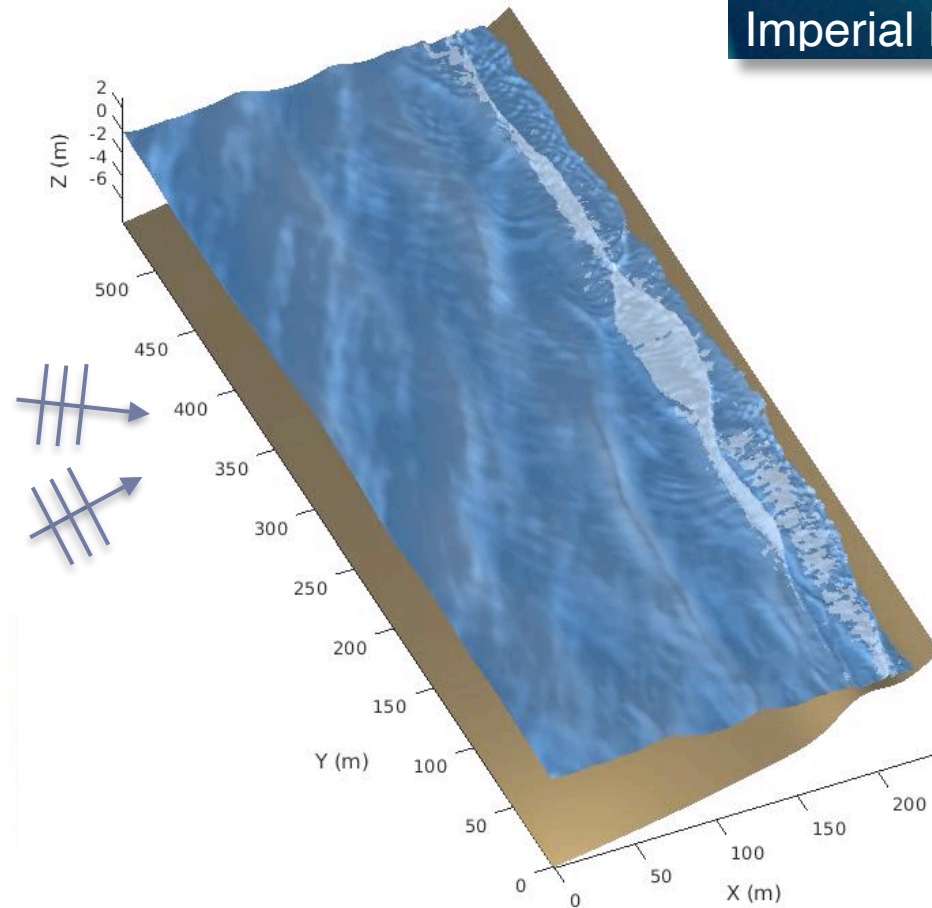
☒ **Turbulence production is limited when strain rate is much greater than vorticity**



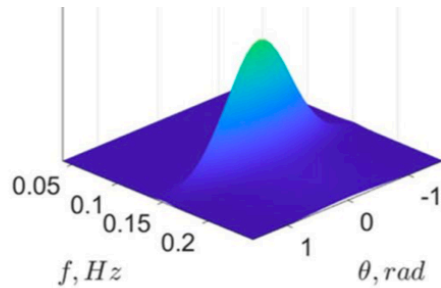


3D Wavemaker

Application to a natural Beach



JONSWAP wave spectrum
with directional spread



Wavemaker correction for the standing wave problem

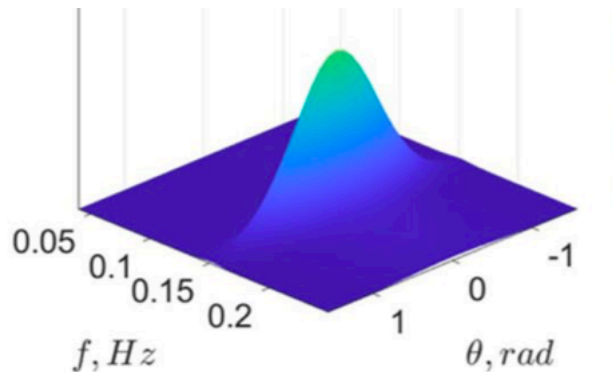
Salatin et al., 2021

Wave-resolving models generally use a classical double summation wave-maker:

$$\eta_{bc}(y, t) = \sum_i a_i \sum_j d_j \cos(k_{y,i,j} y - \omega_i t - \phi_{ij})$$

Amplitude depending on a frequency spectrum
Directional spread

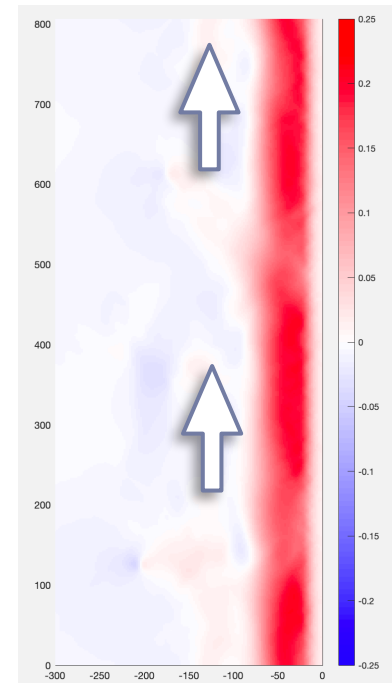
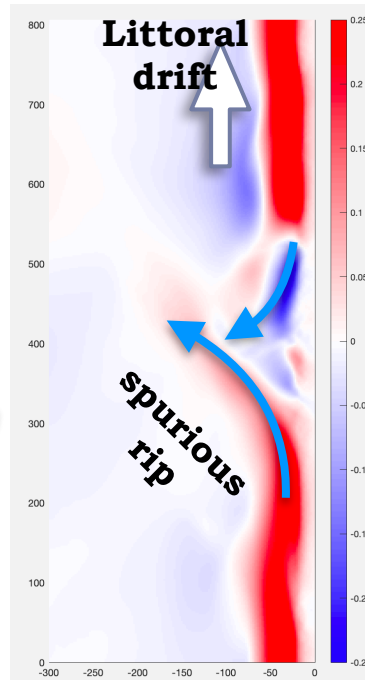
... leading to interferences between waves of different directions with same frequencies (phase-locking)



Waves directional spread

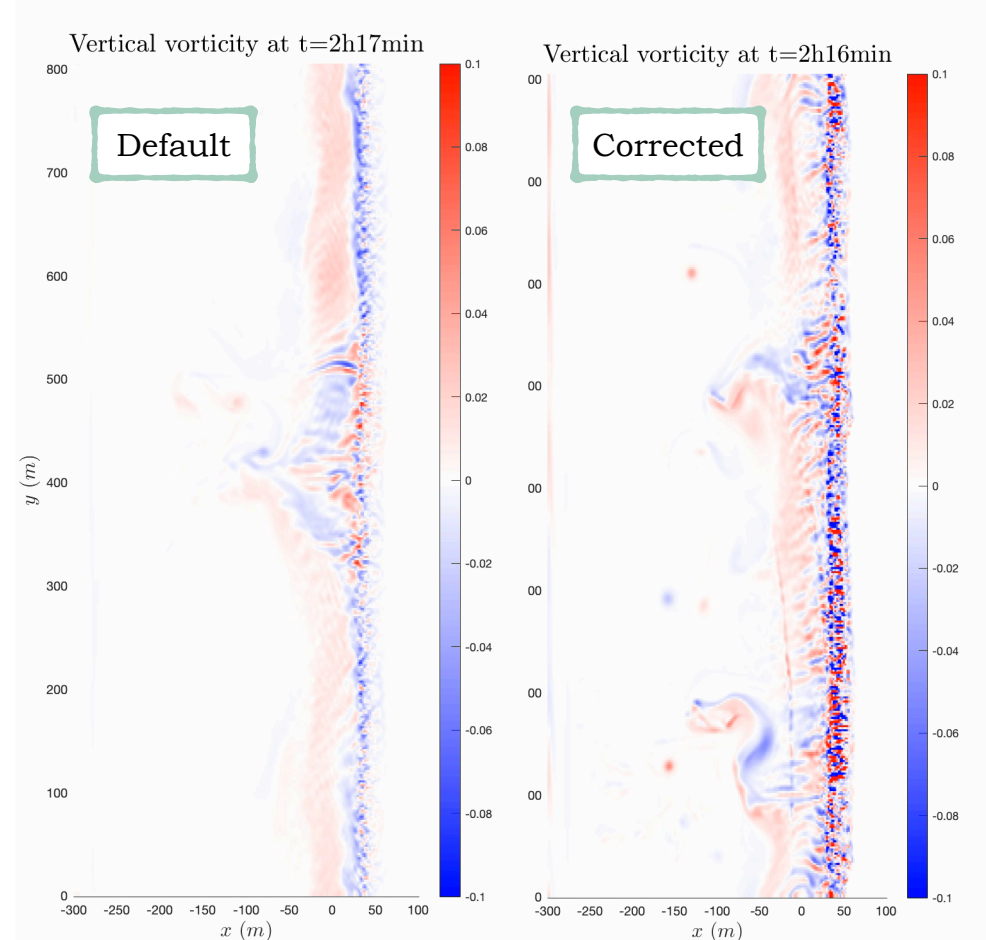
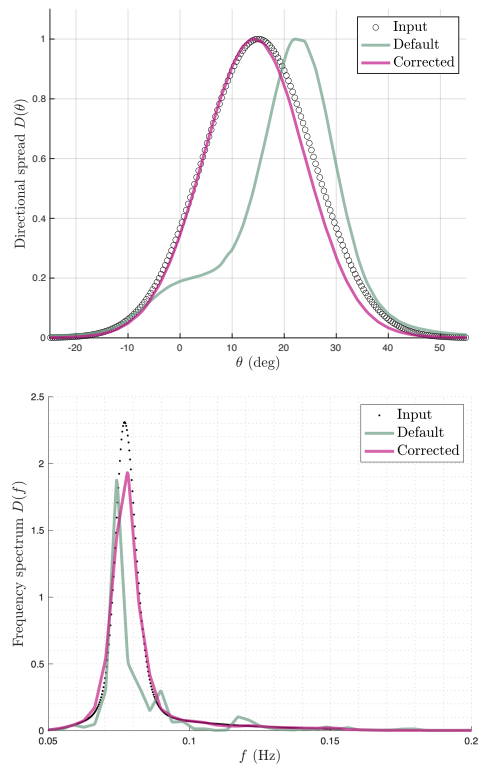
Wave-maker modification such that all wave components have distinct frequencies and directions:

$$\eta_{bc}(y, t) = \sum_i a_i d_i \cos(k_{y,i} y - \omega_i t - \phi_i)$$



Wavemaker correction for the standing wave problem

Modification of the wave-maker such that all wave components have distinct frequencies and directions



☒ **Avoid negative impact on rip currents, cross-shore exchanges, wave streaming...**