

CROCO

Coastal and Regional Ocean COmmunity model

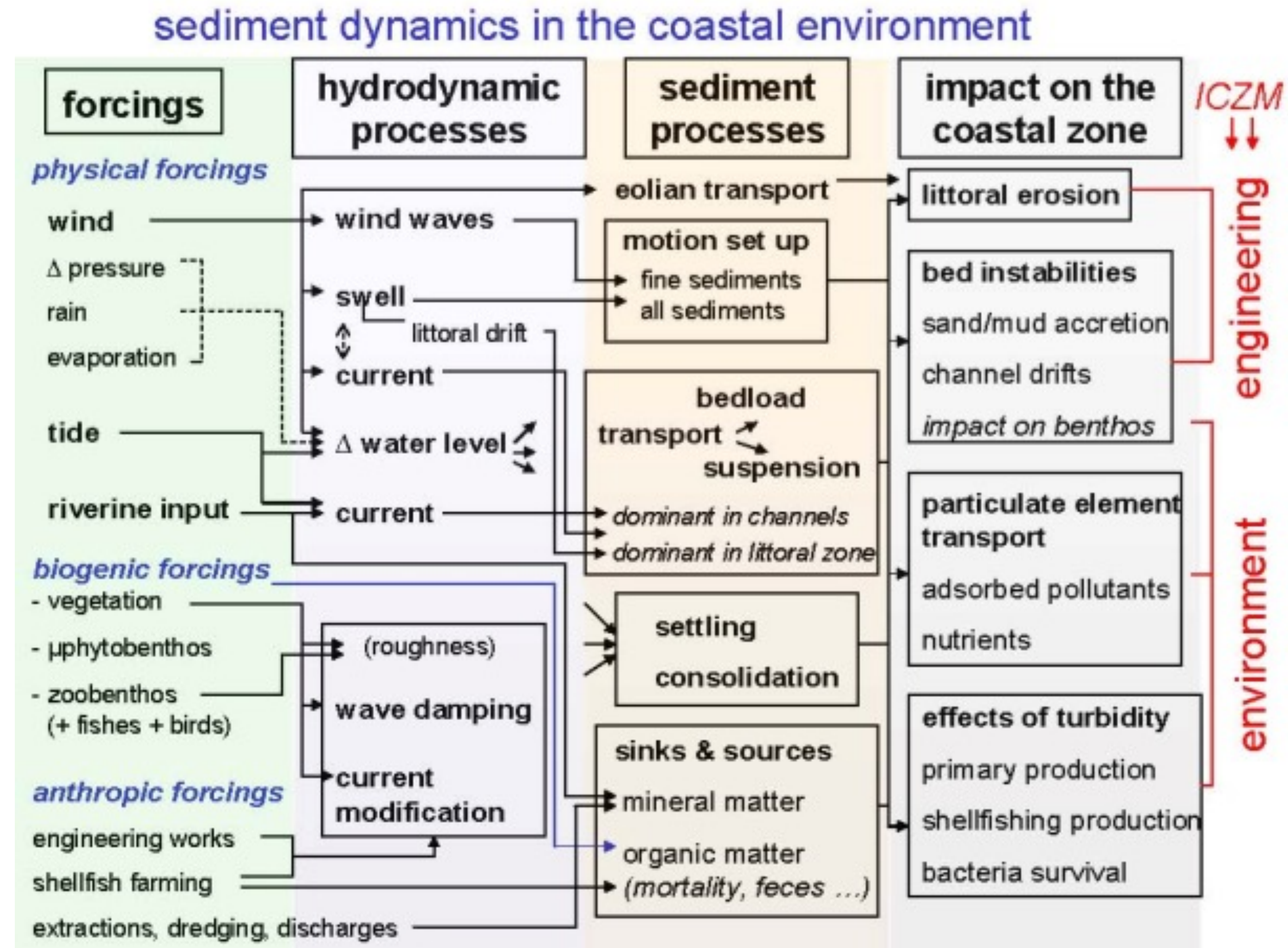
Sediment dynamics

Introduction to sediment transport

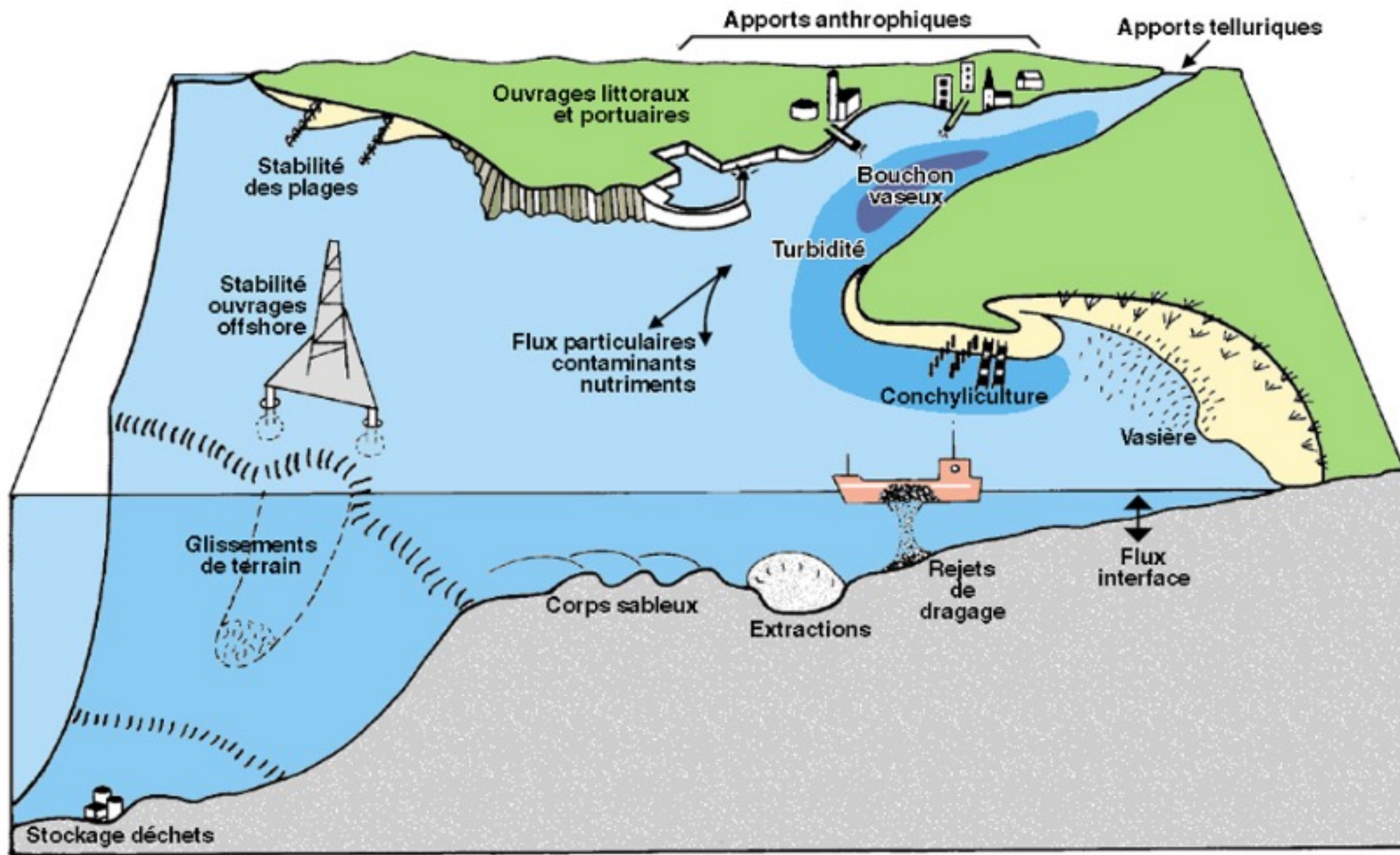
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https://croco-ocean.gitlabpages.inria.fr/croco_doc

Introduction



Introduction

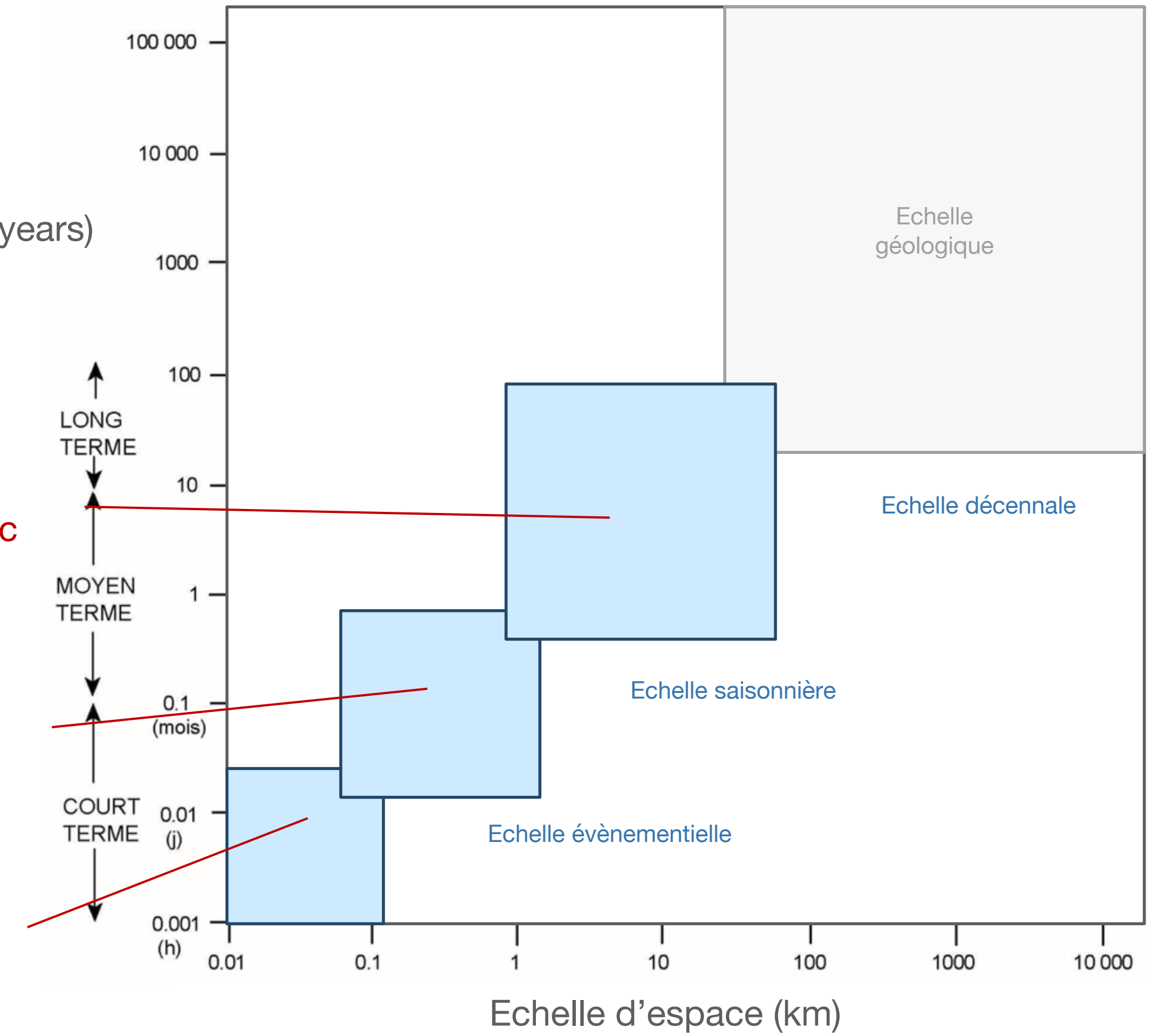


Time scale (years)

Estuarine morphodynamic

Sedimentary fluxes

Sedimentary processes



D'après Fenster et al. [1993] et Dehouck [2006]

Outline

**Forcing drivers : roughness and bed shear stress,
currents and waves**

Sediment properties and behaviors

Modes of sediment transport

In general

Tidal estuary mud transport

Wave driven transport near shore

Modelling strategies

Forcing

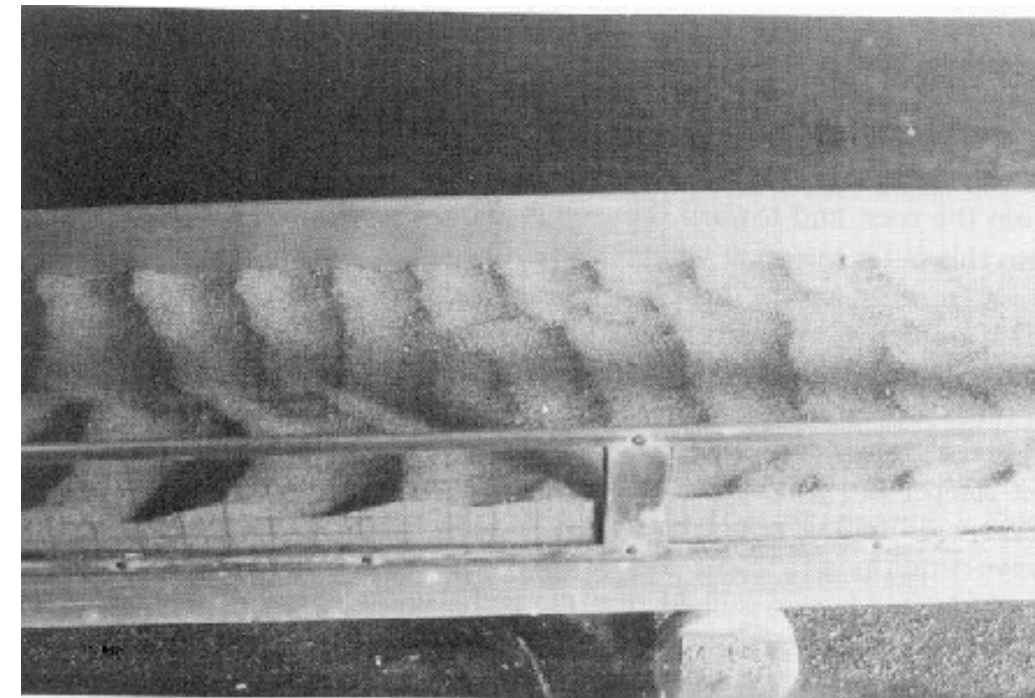
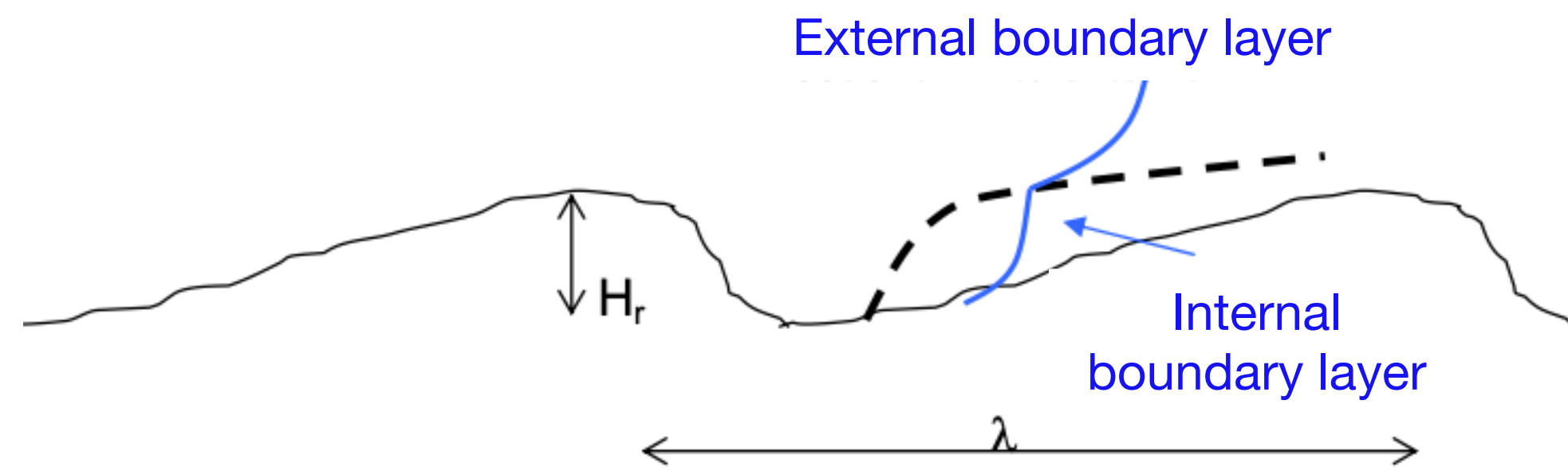
Roughness length - distinguish grain or form related

Depends on sea bed nature : smooth, rough, rippled

Grain related : Nikuradse formula $z_0 = k_s/30$ with $k_s = 2.5D$

Form related :

- ripple prediction formulation from current/wave shear stress
- roughness with ripple bed (Grant and Madsen 1982, Wiberg and Rubin 1988, Van Rijn 1989)



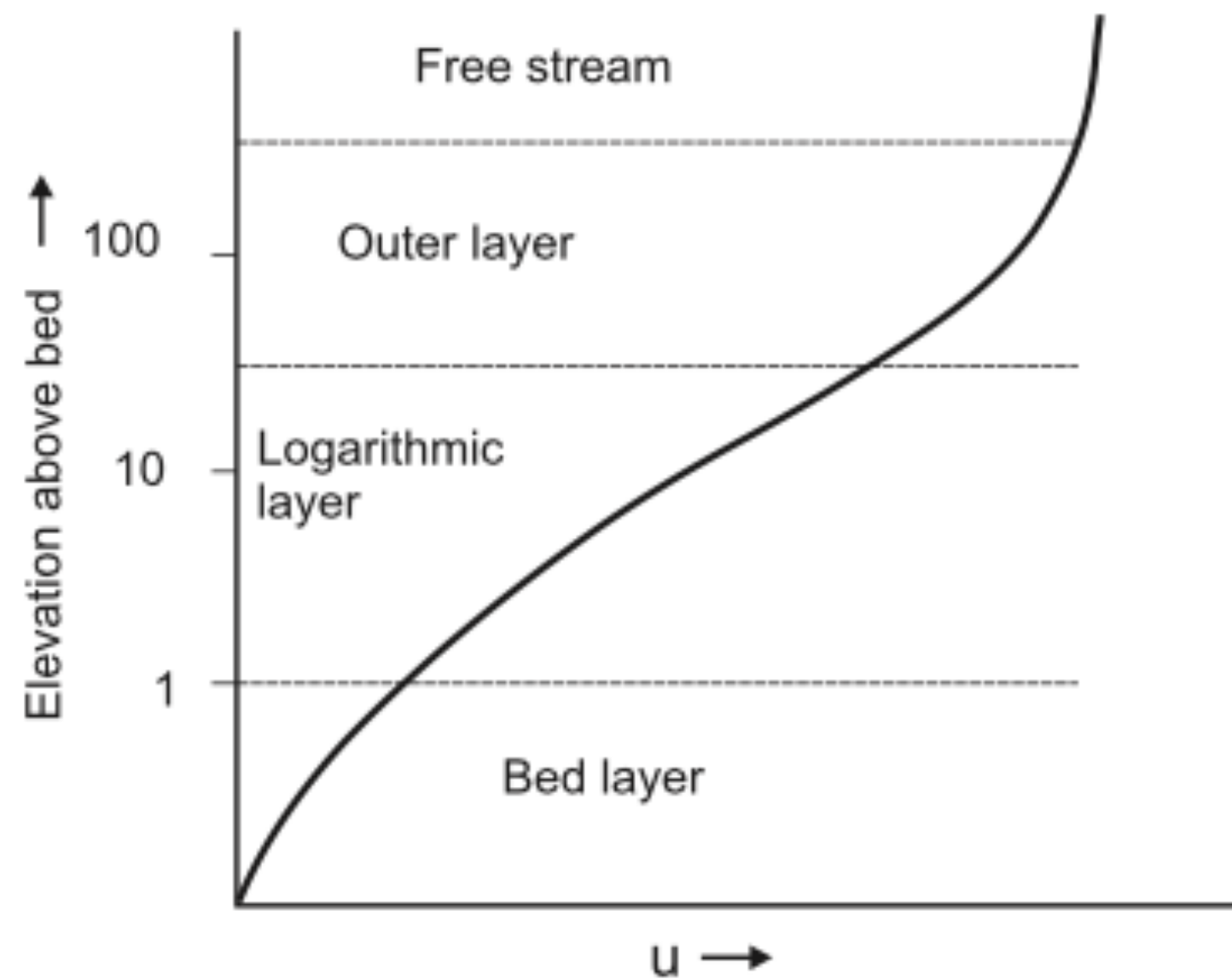
Forcing

Current and wave shear stress

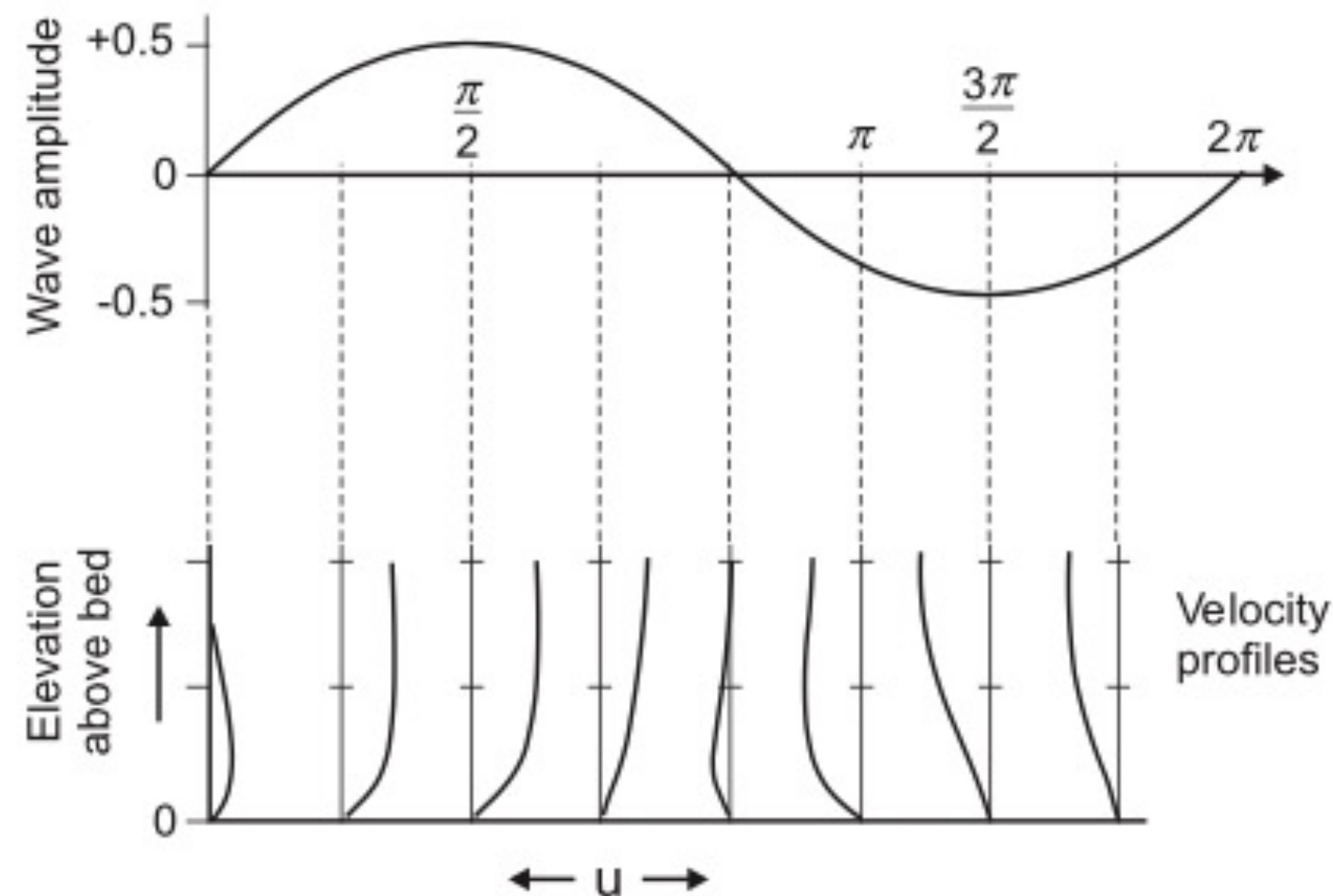
- account for current stress at the bottom
- account for wave shear stress
- sediment dependent z_0

$$\bar{\tau}_{wc} = \tau_c \left(1 + 1.2 \left(\frac{\tau_w}{\tau_w + \tau_c} \right)^{3.2} \right)$$

$$\tau_c = \frac{\kappa^2}{\ln^2(z/z_0)} |u|^2$$
$$\tau_w = 0.5 \rho f_w u_b^2$$



Unidirectionnel current
(tide, rip, wind)



Single wave

Forcing

Summary

Use skin friction with roughness length depending on bed composition

$Z_{0sed} \neq Z_{0hydro}$

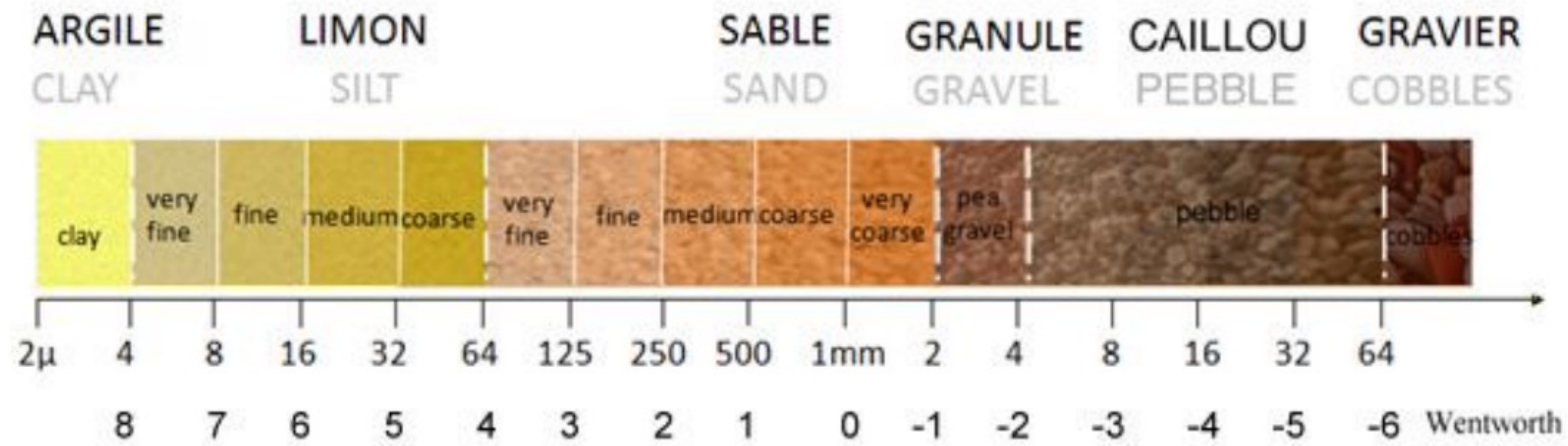
Interaction between current and wave : combination of shear stresses

Needs to know hydrodynamic variables :

**depth, current, wave period, wave orbital velocity,
water density, bed roughness length**

Sediment properties

Grain size



Sediment can be characterized by grain size (d50, d90)

- **cohesive** : mud (clay, silt)

- transport in suspension
- flocculation, variable settling velocities
- bed consolidation

- **non cohesive** : sand to cobble

- transport in suspension and/or bedload

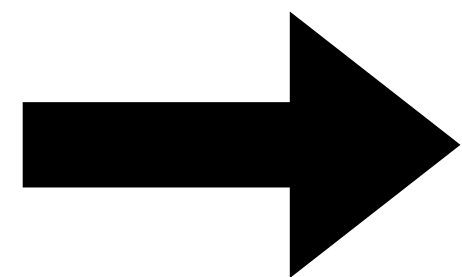
Sediment properties

Settling velocity

Cohesive

Stokes if isolated grains

Links between grains leads to variable settling velocities (flocculation, hindered effect)



Several formulations (Migniot, Thorn & Parsons, Van Leussen) or modeling approaches (Winterwerp, Wolanski)

Non cohesive

No link between grains
Hindered effect with high concentration

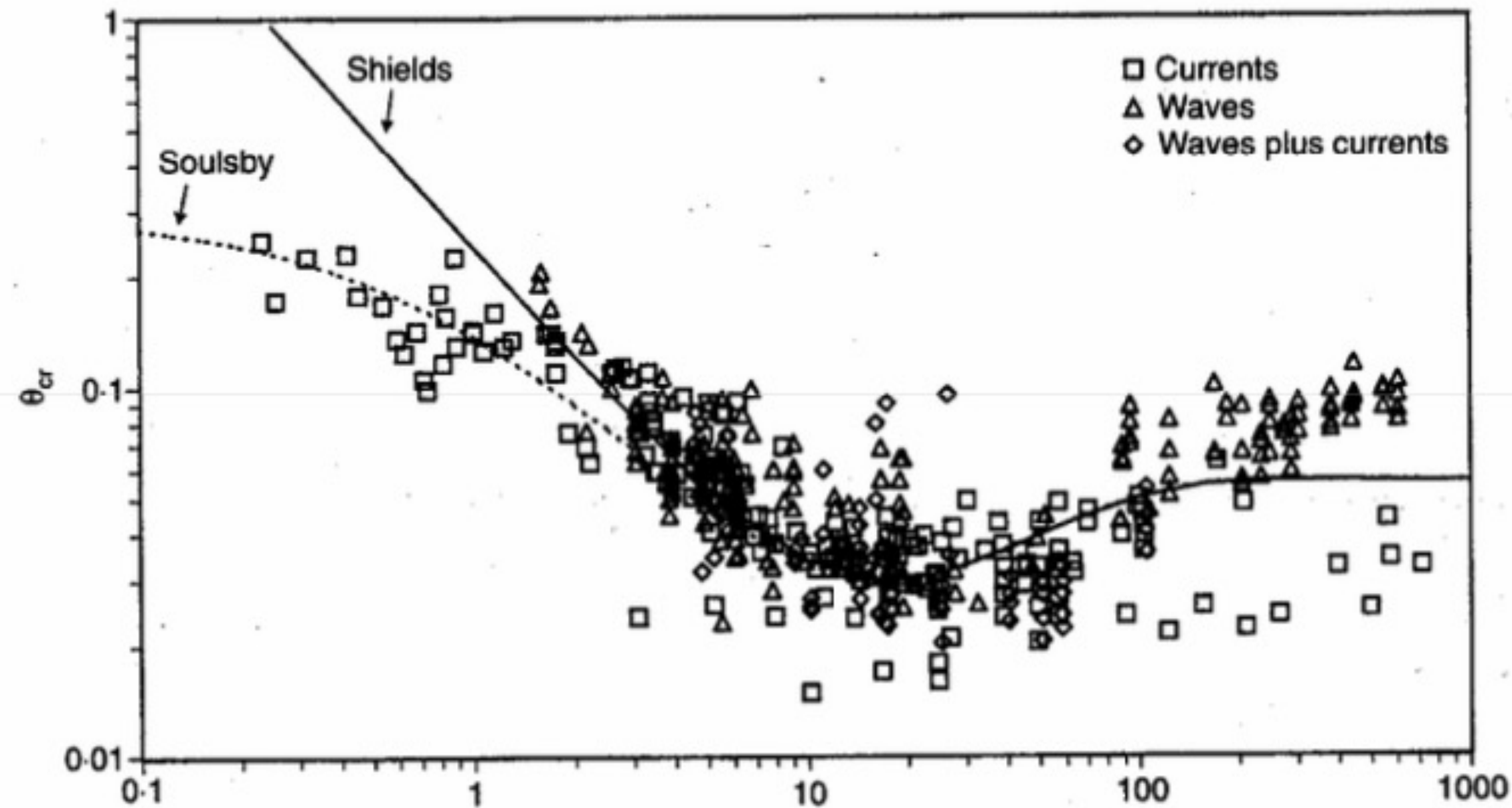
Wide range of values
~1 cm/s for ~100 μm grain size
~10 cm/s for ~800 μm grain size

Soulsby (1997)

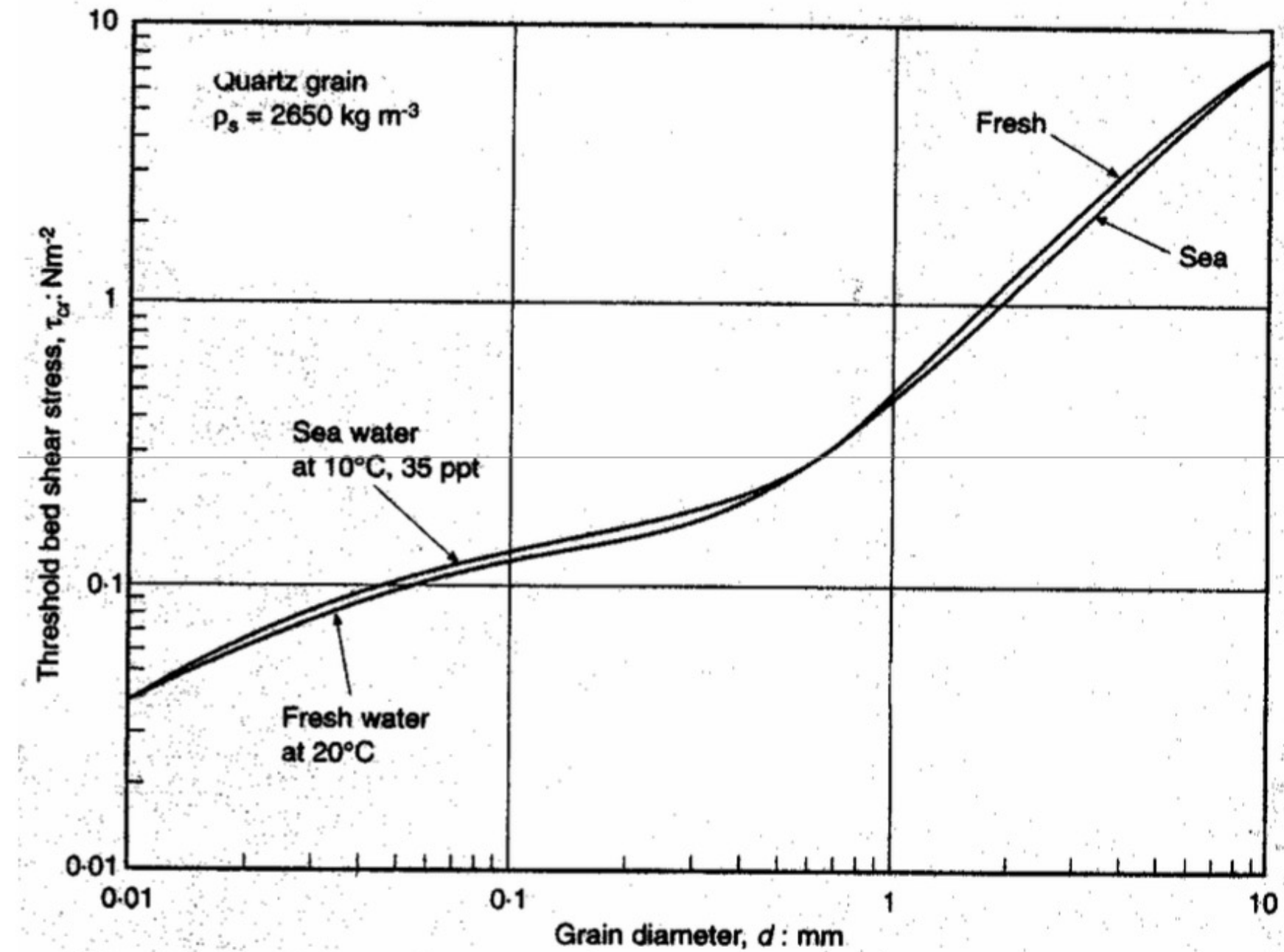
$$W_s = \frac{v}{D} \left[(10.36^2 + 1.049 D_*^3)^{0.5} - 10.36 \right] \quad \text{avec } D_* = D \left[\frac{g(s-1)}{v^2} \right]^{1/3}$$

Sediment properties

Motion threshold - non cohesive case



Soulsby (1997)
$$\theta_{cr} = \frac{0.3}{1+1.2D_*} + 0.055 \left[1 - \exp(-0.02 D_*) \right]$$



Motion threshold - cohesive case

Depends on consolidation and cohesion of the sea bed

Sediment properties

Mixed sediment, multi class : sorting, bed armoring, mask-exposure

**Modification of threshold of motion from sediment granulometry
repartition in sea bed**

Modification of transport capacity (hindering coefficient)

**Modification of erosion fluxes depending on fraction of cohesive
sediment in bed (cohesive, intermediate or non-cohesive)**

Modification of porosity (deposition of fine sediment in coarse bed)

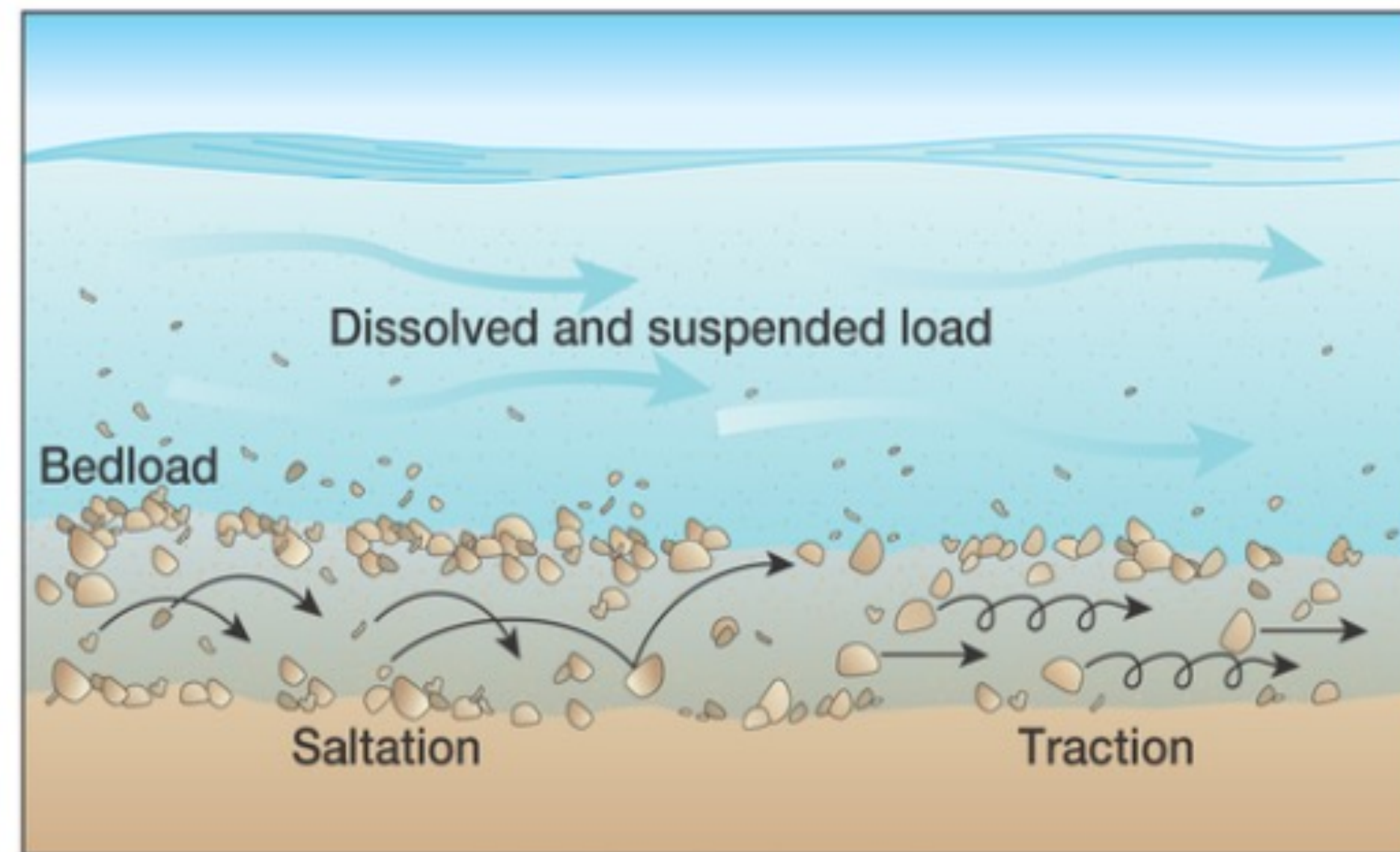
Modes of sediment transport

BEDLOAD

Transport in contact with the bed :
rolling, saltation

SUSPENSION

Transport in water



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All simultaneously

Modes of sediment transport

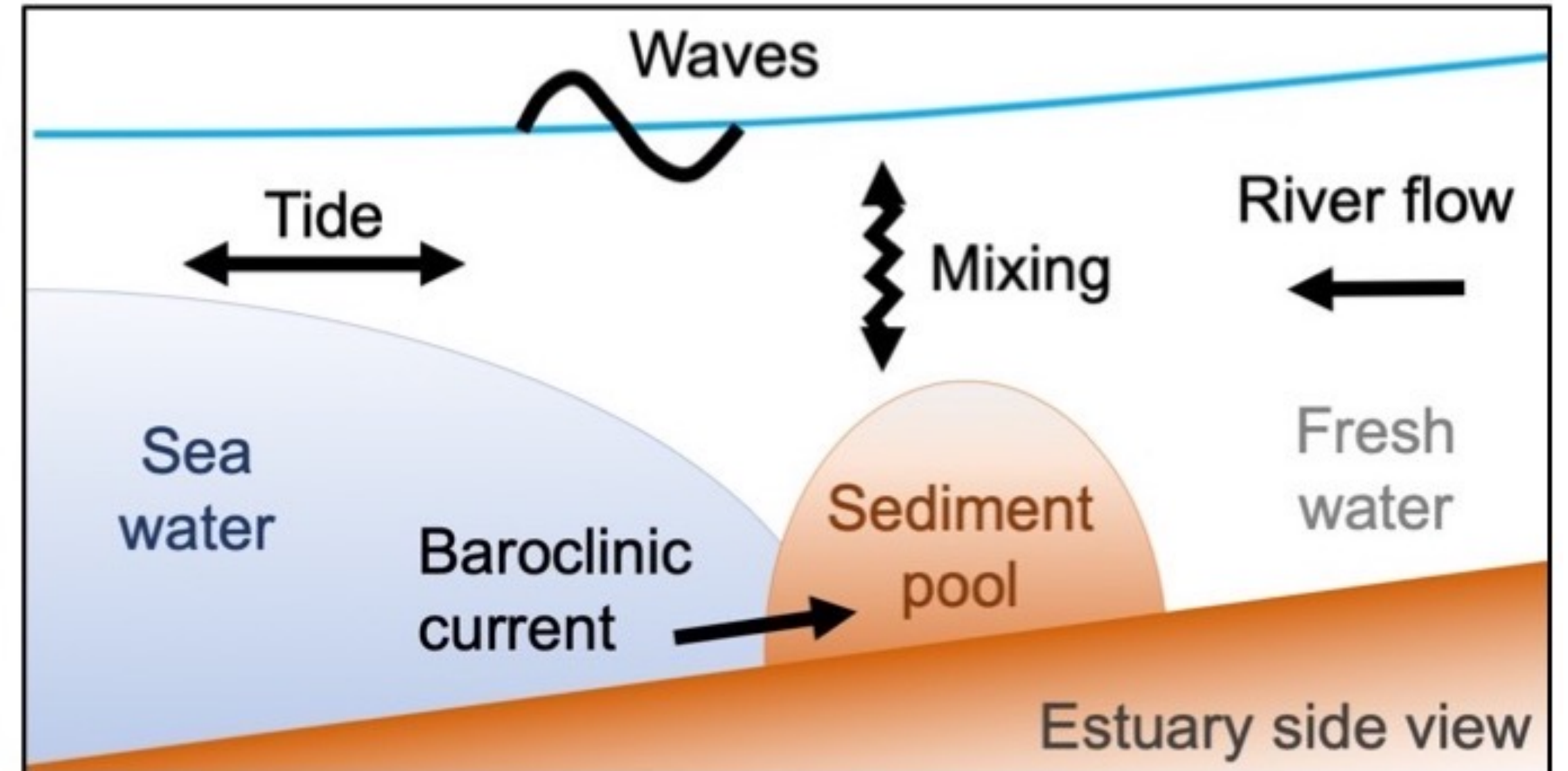
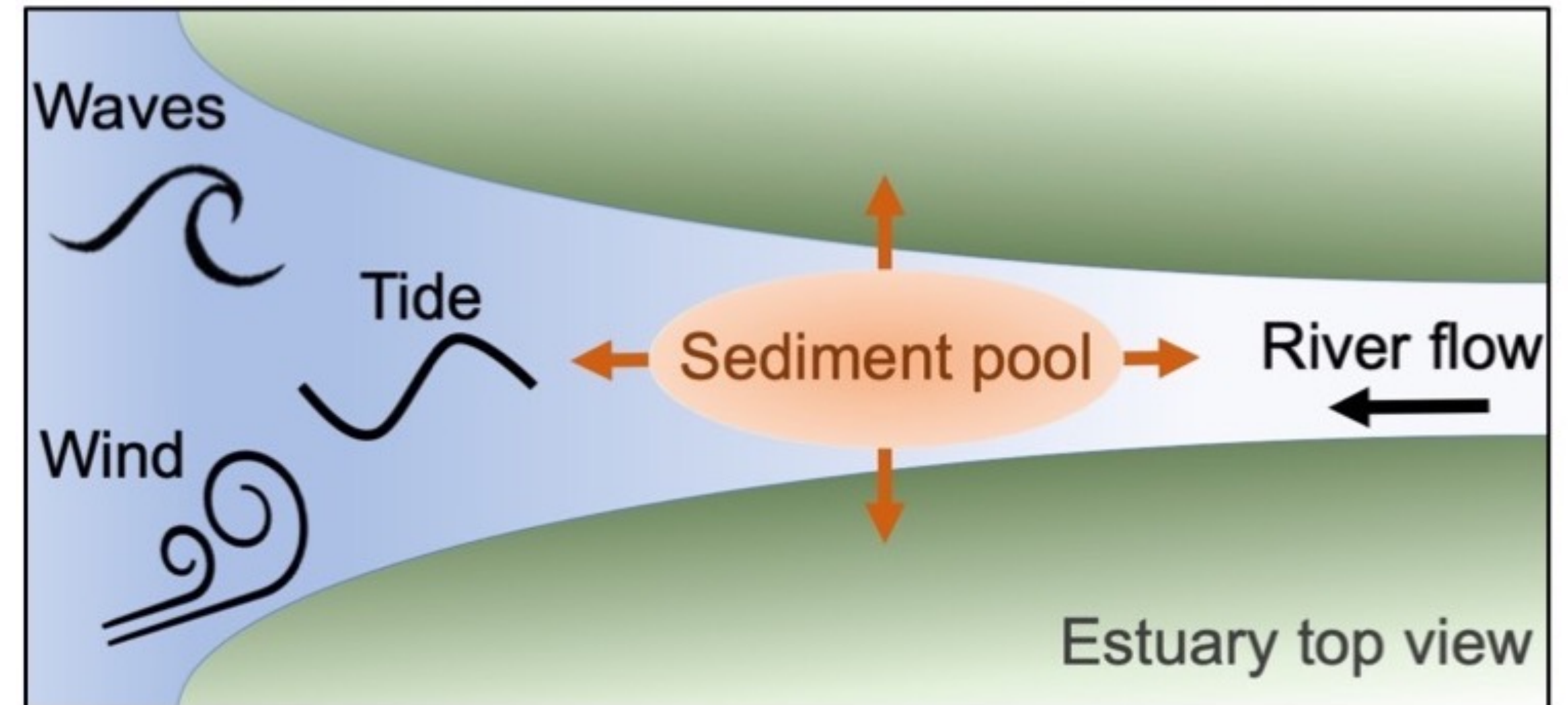
Tidal estuary mud transport

- Forcings

- Oceanic (tide)
- Meteorological (wind, waves, surges)
- Fluvial (liquid and solid flows)

- Area of high turbidity

- Suspension of sediments
- Sediment convergence

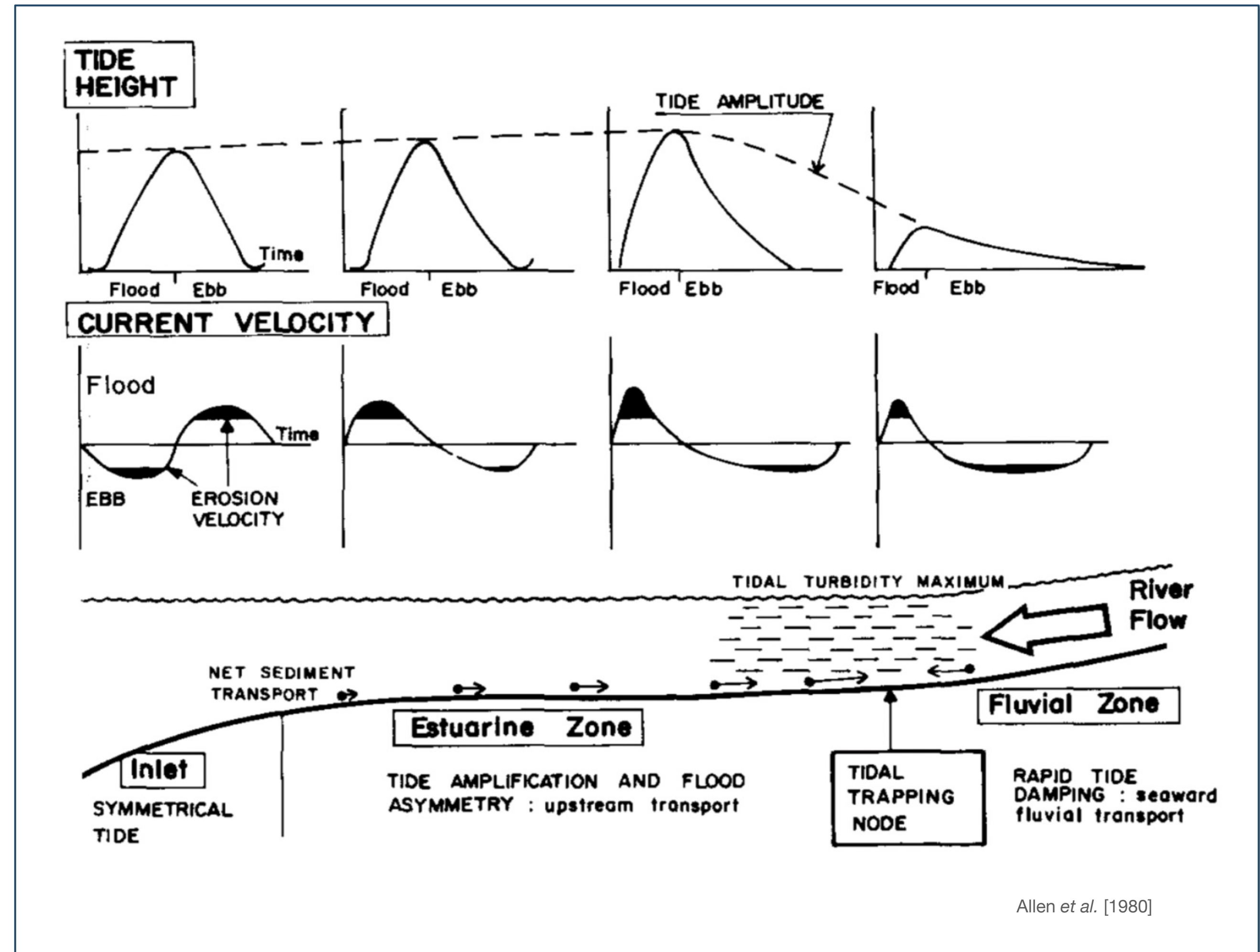


From Burchard *et al.* [2018]

Modes of sediment transport

Tidal estuary mud transport

- River flow
 - Downstream transport
- Tidal asymmetry
 - Tidal pumping
 - Upstream transport
- Baroclinic circulation
 - Difference in density : fresh water / salt water
 - Upstream transport



Modes of sediment transport

Wave driven transport near shore

We distinguish :

- processes that lead to net transport of sediment onshore or offshore (cross-shore transport)
 - processes tending to move sediment alongshore (longshore sediment transport)
- => Both occurs simultaneously

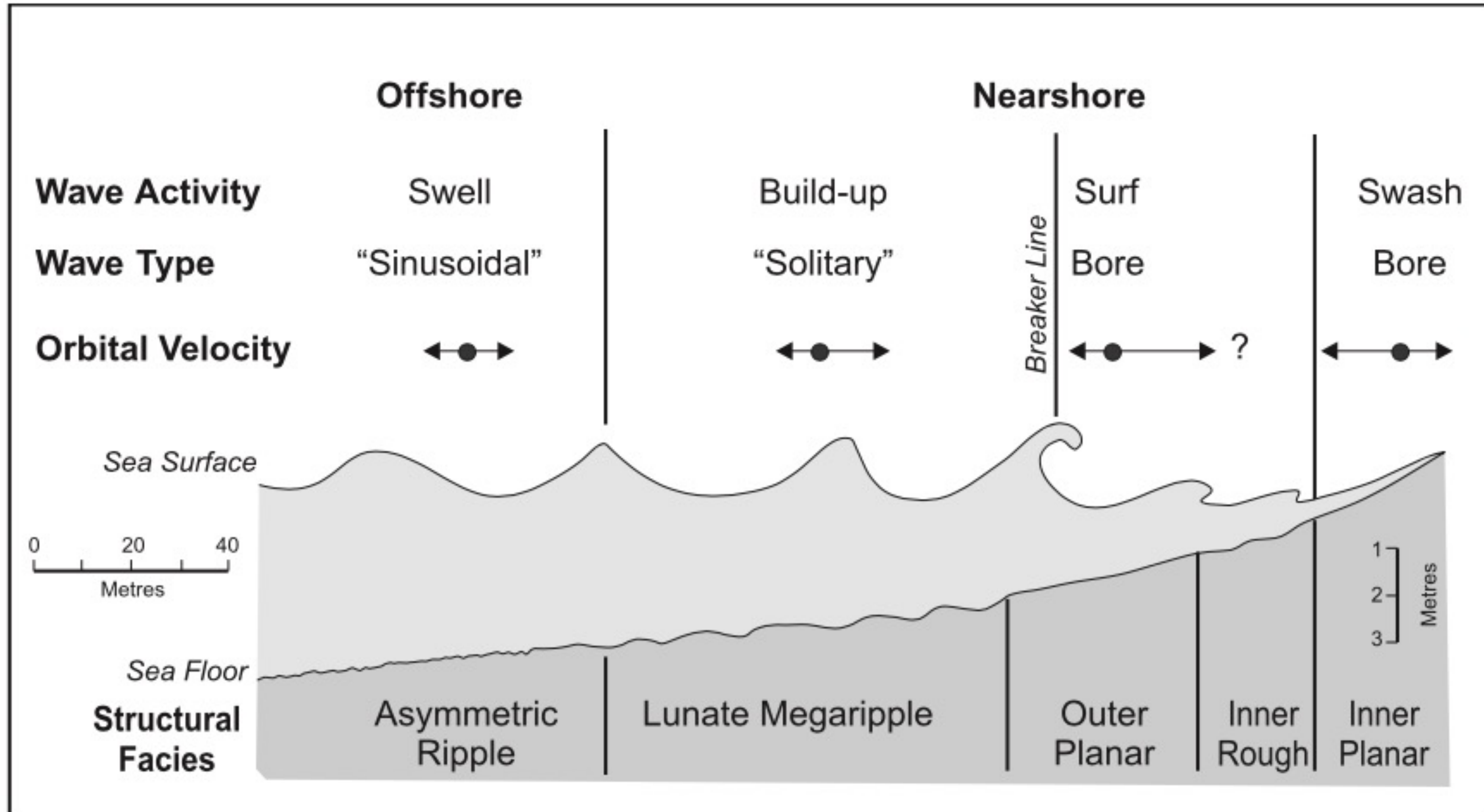
Set of motion :

- generally no erosion and transport by unidirectional currents (except RIP, strong longshore drift)
- Instead :
 - motion by oscillatory currents due to waves and wave-breaking turbulence
 - transport by mean flow : undertow, stokes drive, wind-drive

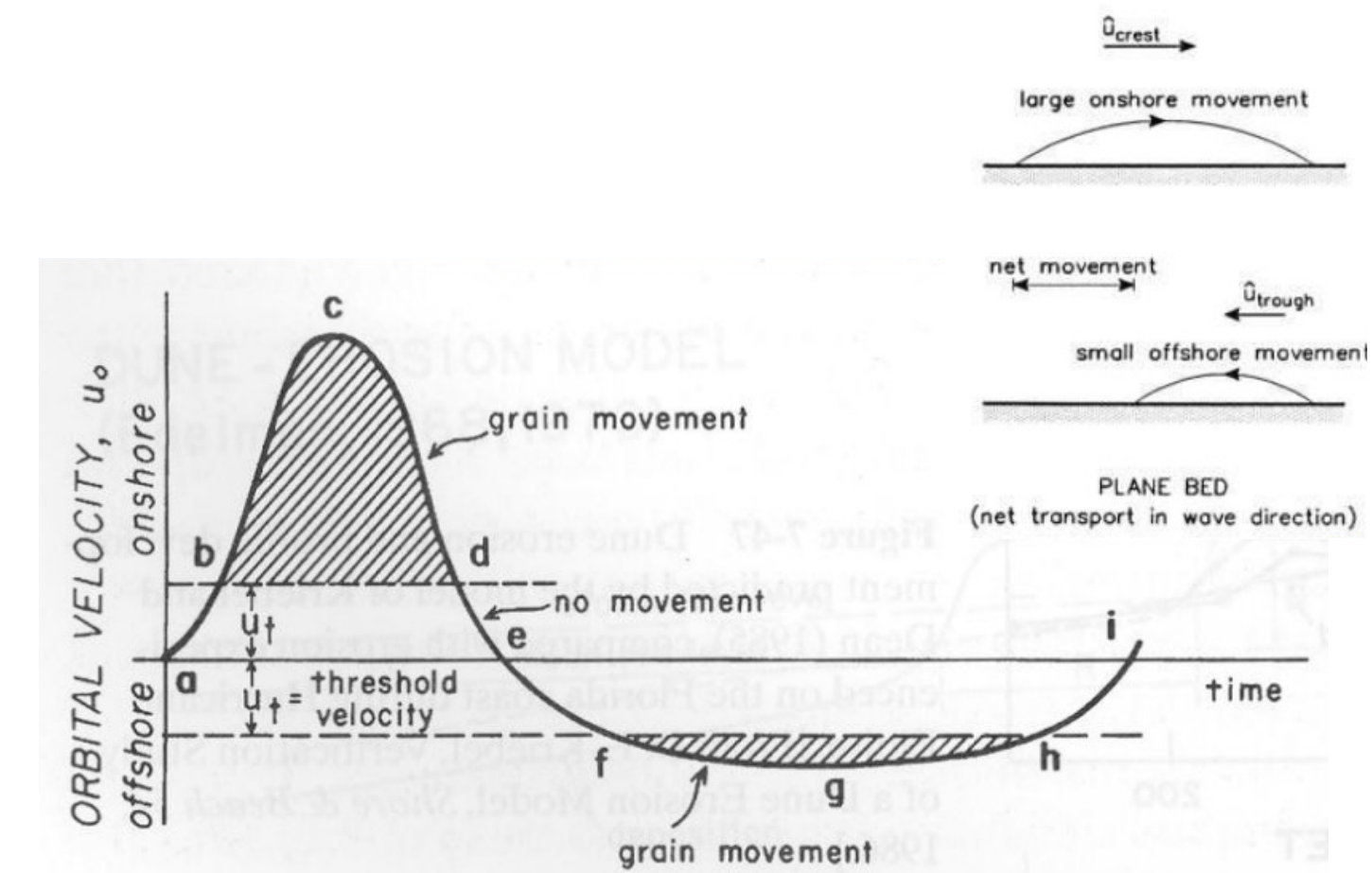
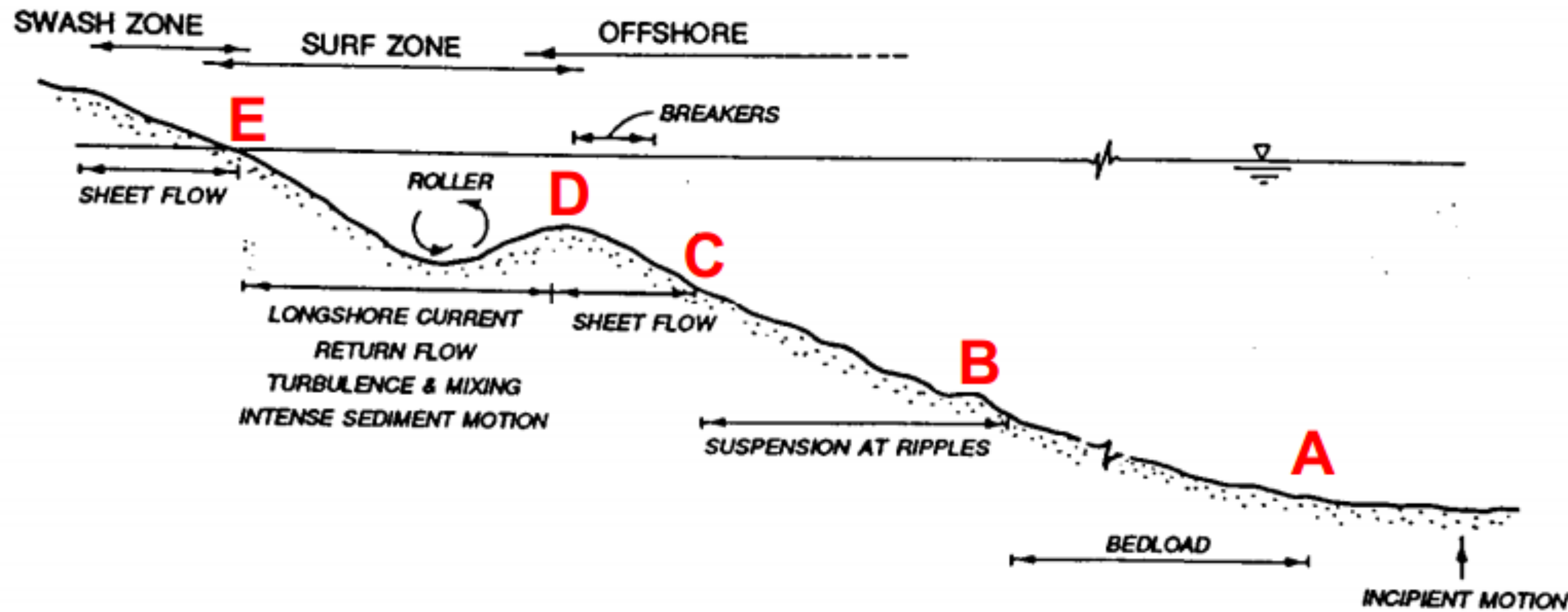
Direction of net transport of sediment : the balance of all

- incident and wave-generated on/offshore and alongshore flows
- wind-driven currents
- tidal flows

Waves at nearshore

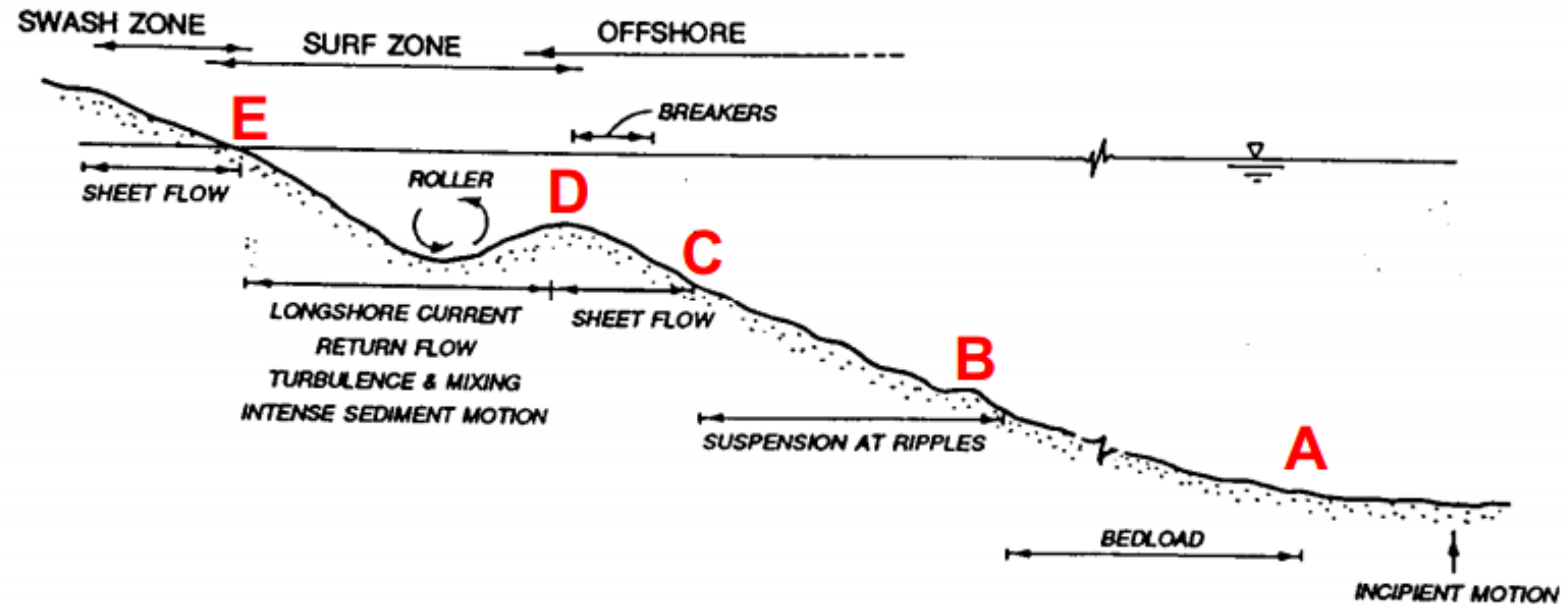


Cross-shore transport



- A : no transport (symmetric waves, inactive bed)
 - B : little transport (skew waves and ripples)
 - C : onshore transport in shoaling zone (skewed waves)
 - D : on/offshore transport in break-in zone (asym waves / undertow)
 - E : on/offshore transport in swash zone (infra gravity waves)
- => transport rate increase

Cross-shore transport



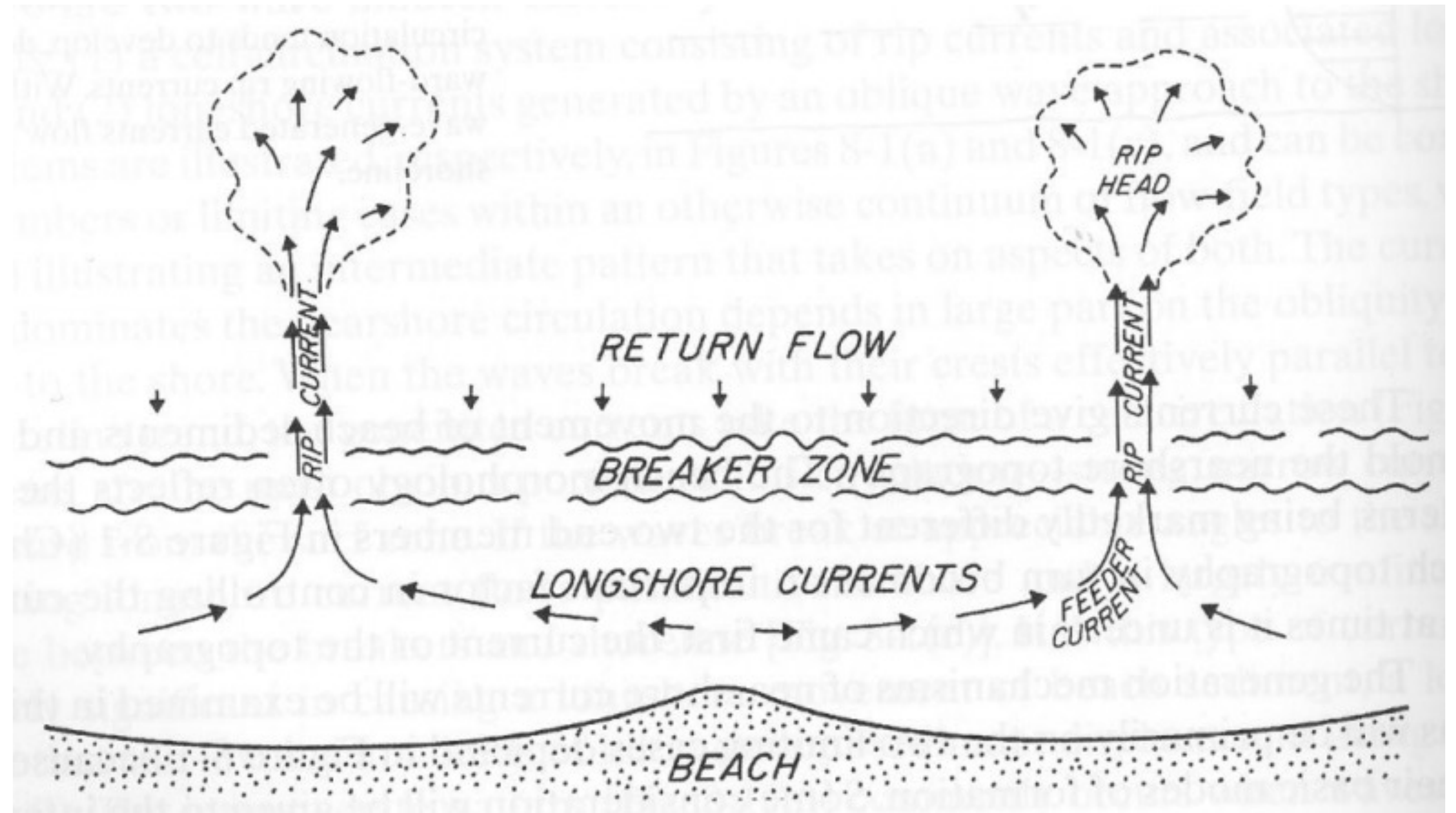
Toward long term :

- Locations A-D shift as a function of offshore wave conditions
- A specific cross-shore location experiences many different conditions during a year
- Which are dominant :
 - Frequent low energy conditions ?
 - Occasional storm ?

Cross-shore transport

Particular case of rip currents :

- Sediment stirred by gravity waves, transported by currents
- Minor role (onshore) in between the rip currents
- Other mechanisms minor (no undertow!)



Alongshore transport

Gravity waves stir sediment

Breaking induced alongshore current transport sediment

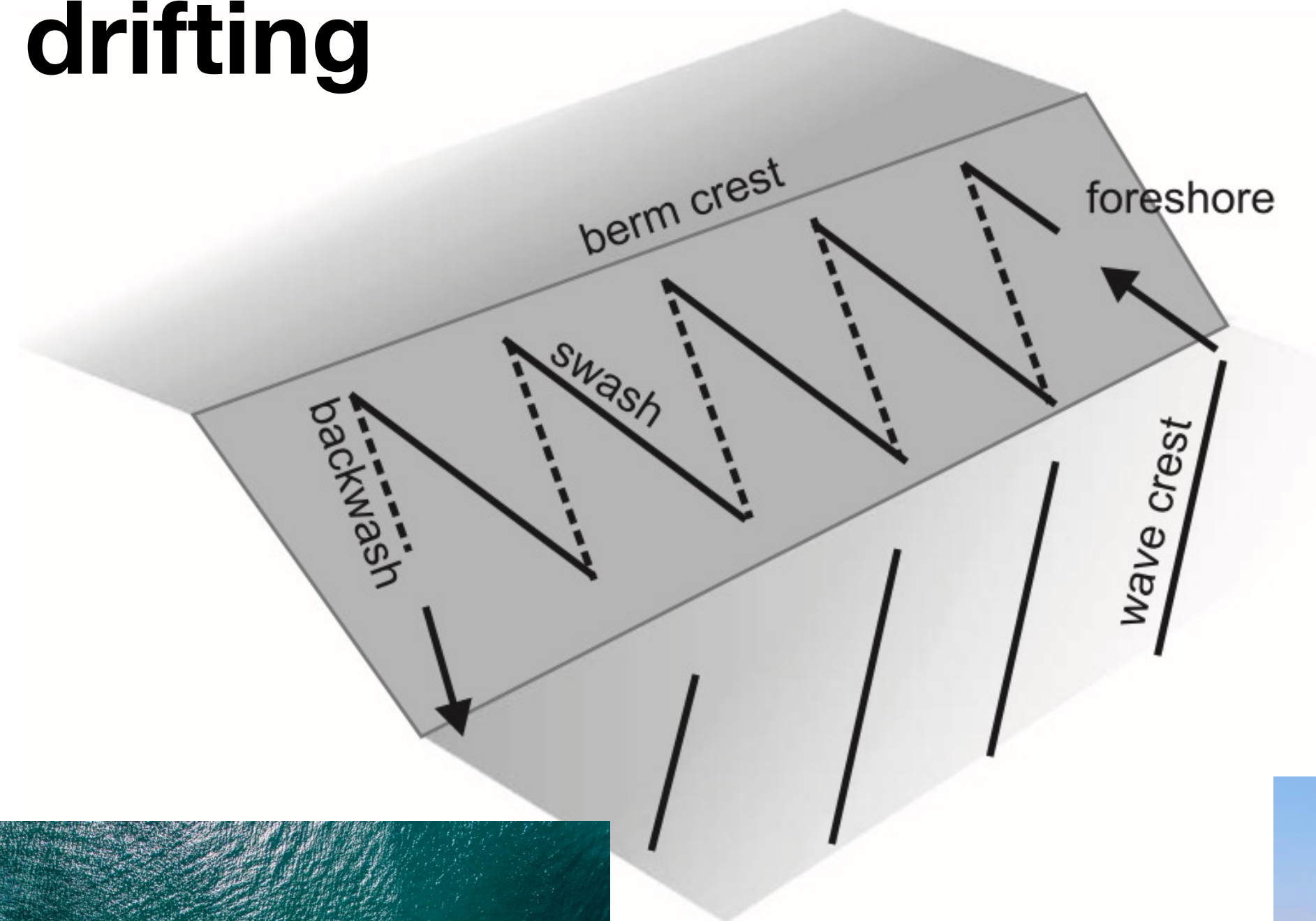
=> Littoral drift

Drivers of sediment transport in the beach and nearshore zone :

- beach drifting on the swash slope driven primarily by oblique wave action
- transport by wave-generated longshore currents in the surf zone
- transport seaward of the breaker zone by residual tidal currents and wind-driven currents.

Alongshore transport

Beach drifting



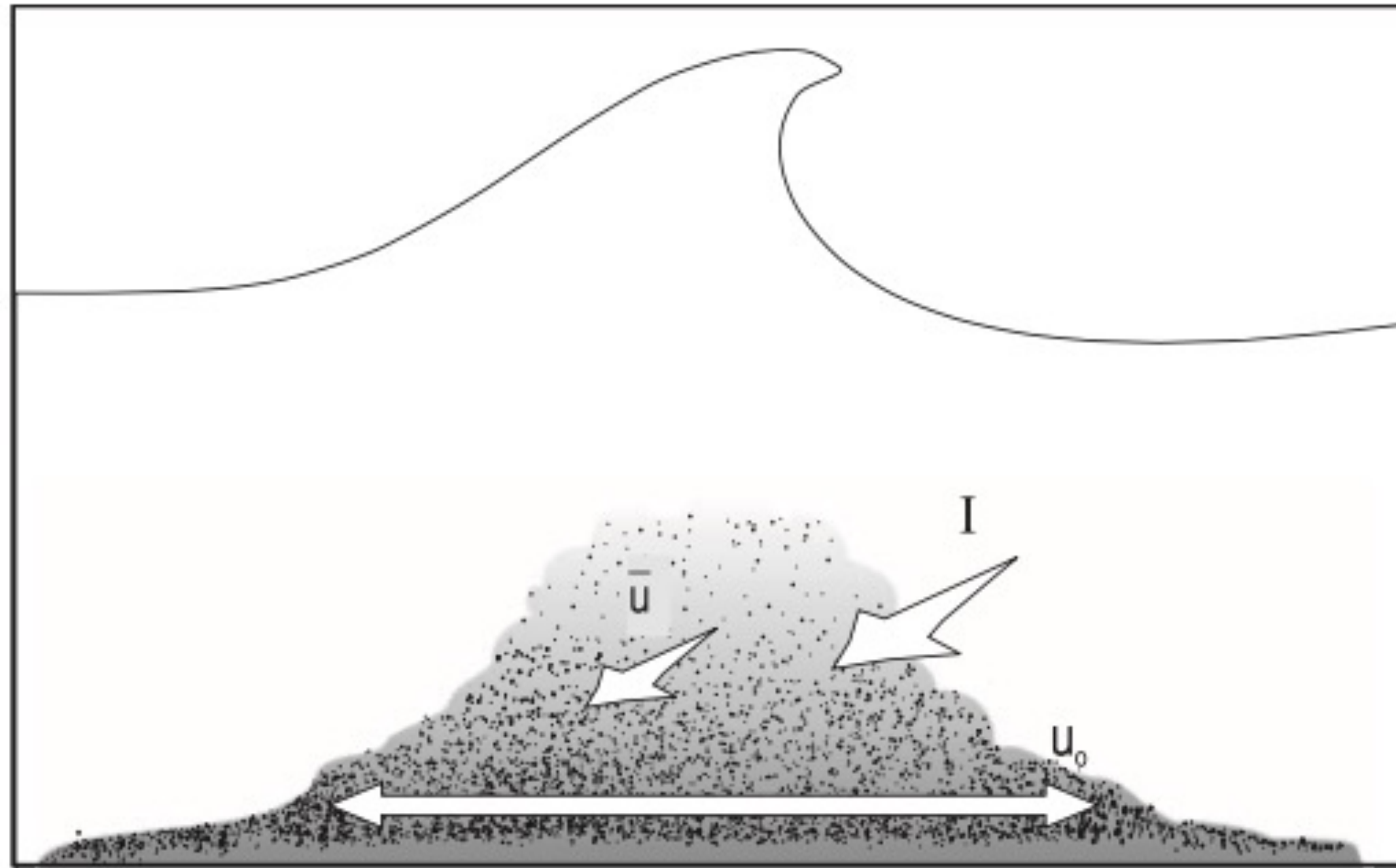
- Swash run-up perpendicular to the wave crest
- Return flow in the backwash occurs parallel to the beach slope (gravity)

=> saw-tooth alongshore motion



Alongshore transport

Surf zone transport : summary



- no transport from oscillatory wave motion
- sediment motion set by wave motion and breaking induced current
- alongshore currents generated in the surf zone by waves breaking at an angle to the shoreline
- wind and tides currents

Alongshore transport

Examples



Alongshore transport

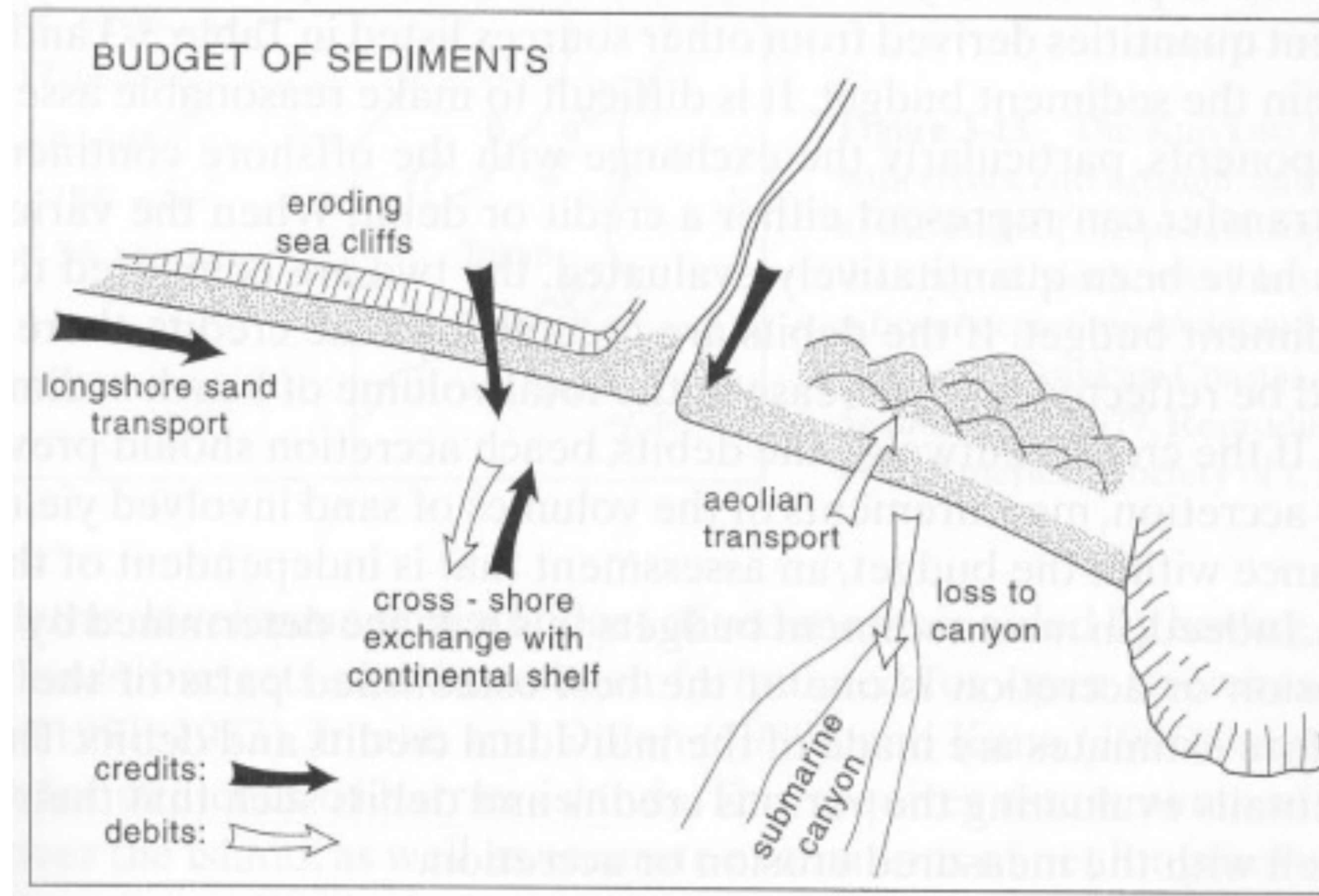
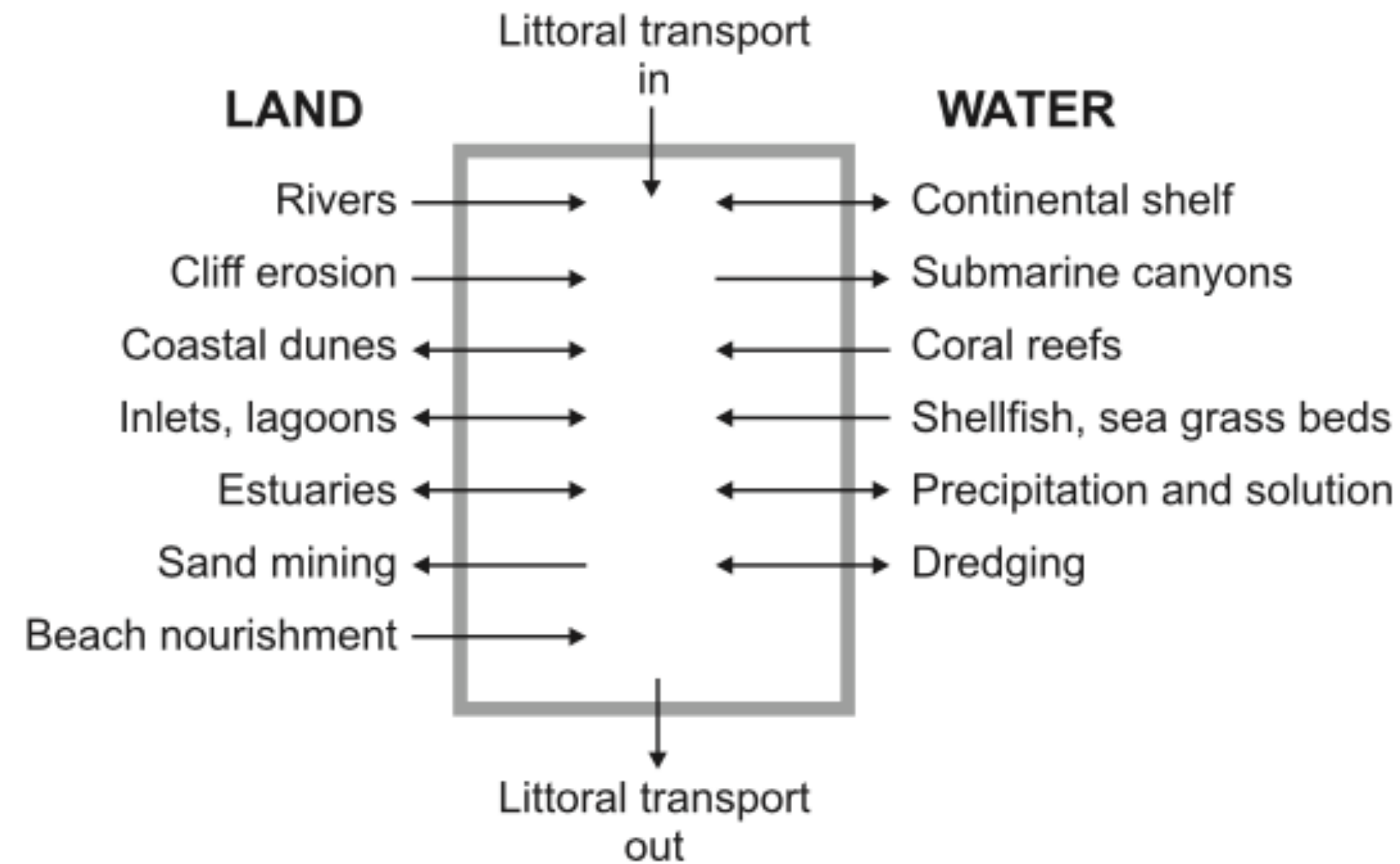
Prediction

Longshore transport empirical formulae :

Correlated with longshore wave energy flux : $q \sim \rho g^{1.5} H_b^{2.5} \sin \theta_b \cos \theta_b$

- Shore normally incident ($\theta_b=0^\circ$ transport is 0)
- Transport increases when wave height increases
- Transport is maximum for $\theta_b = 45^\circ$

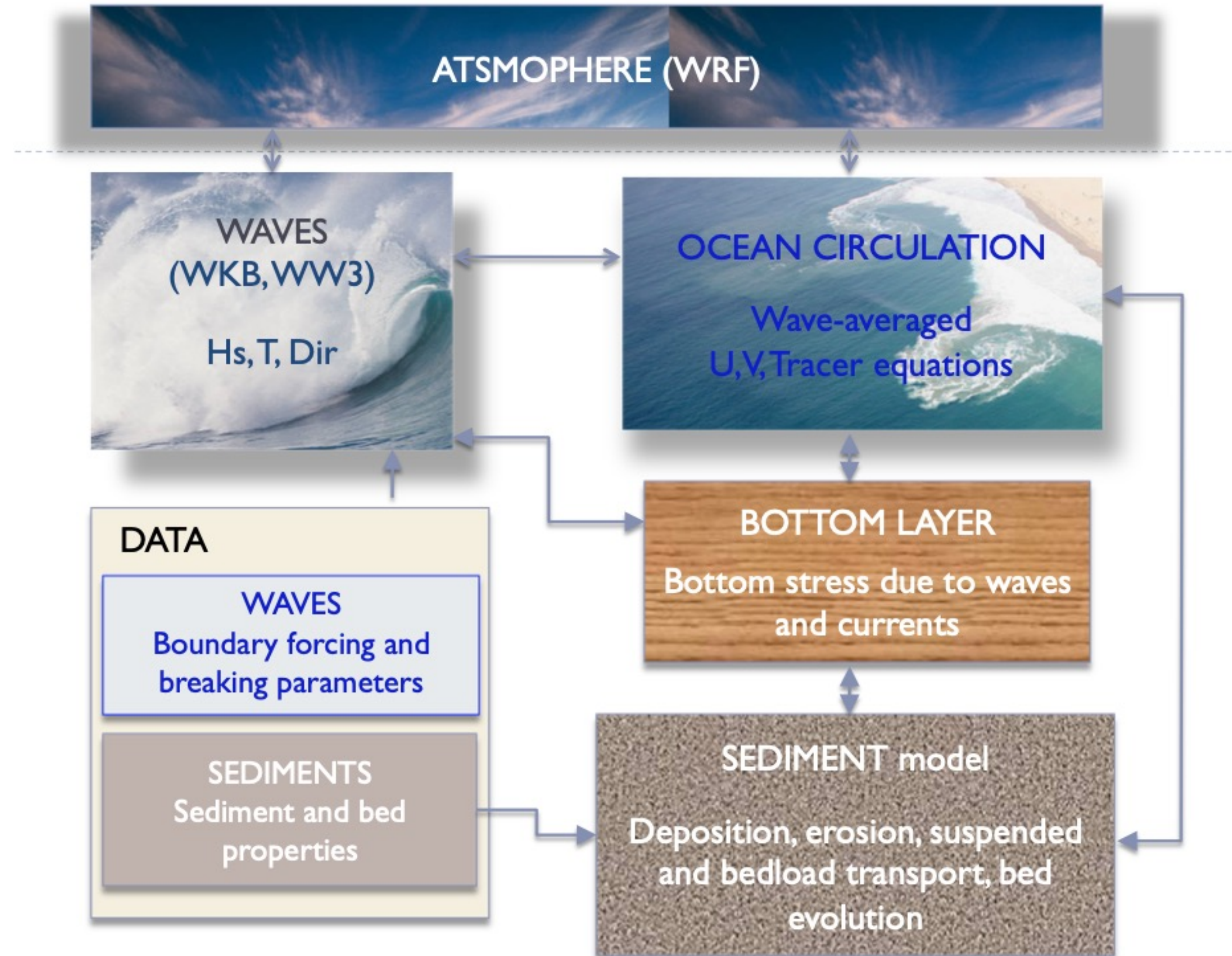
Sediment balance



- bed levels changes are a result of **gradients** in the sediment transport rates
- mass balance equation (Exner equation)

$$\frac{\partial q_x}{\partial x} + \frac{\partial q_y}{\partial y} = \frac{\partial z}{\partial t}$$

Modelling strategies



Modelling strategies

- Transport at equilibrium

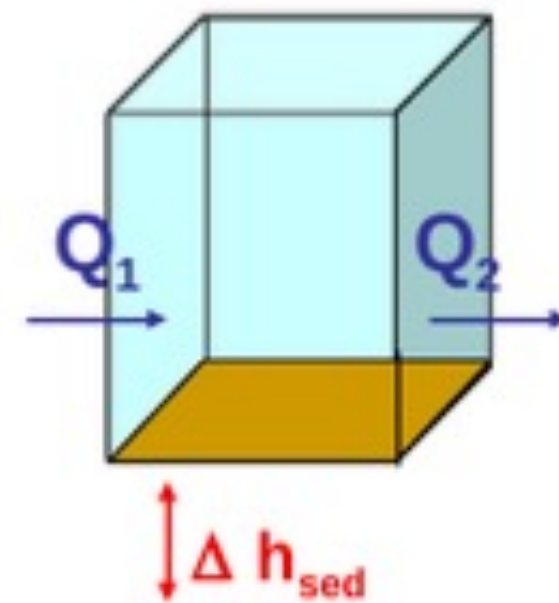
A lot of formulas available :

Bedload, suspension load, total load

Current only, current and wave

Wave asymmetry

+ Slope effect



Transport at equilibrium :
« transport capacity »

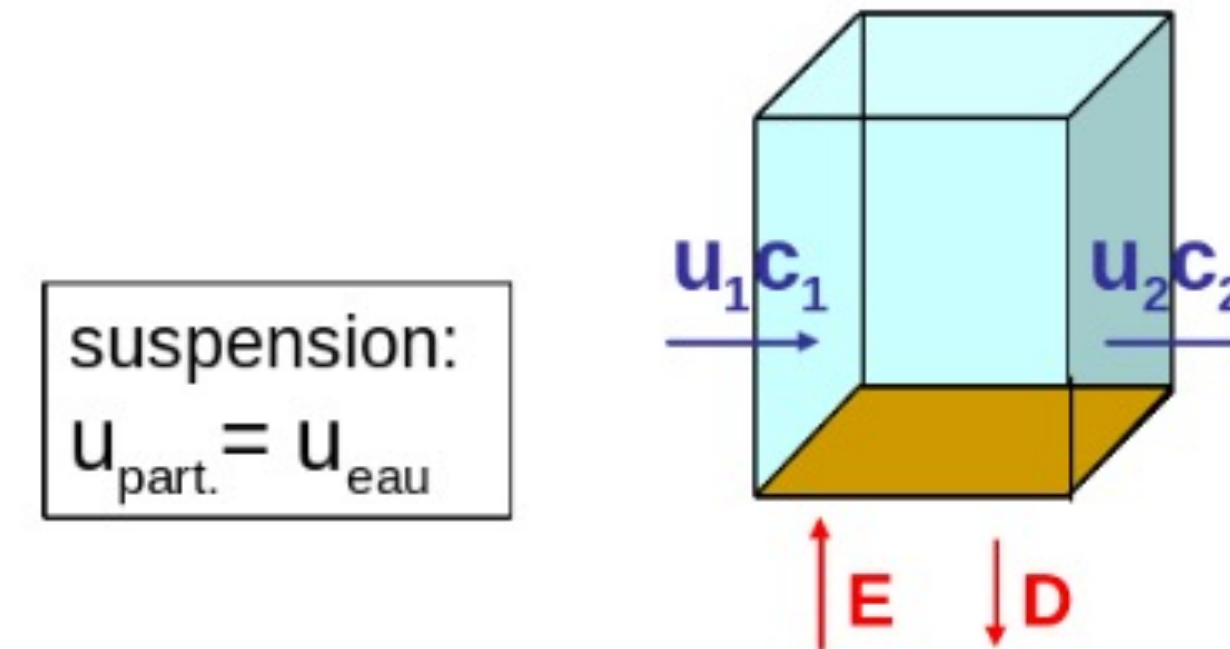
$$Q = f(\text{cond. Hydro.}, \text{param. Sed.})$$

Bottom evolution:

$$(1-p) \frac{\partial h_{sed}}{\partial t} = \frac{\partial Q}{\partial x}$$

- Non-equilibrium transport

Erosion-deposition fluxes



suspension:
 $u_{part.} = u_{eau}$

Erosion flux : $E = f(\text{cond. Hydro.}, \text{param. Sed.})$

Deposition flux: $D = W_s C (1 - \frac{\tau}{\tau_d})$

Eq. Advection/dispersion:

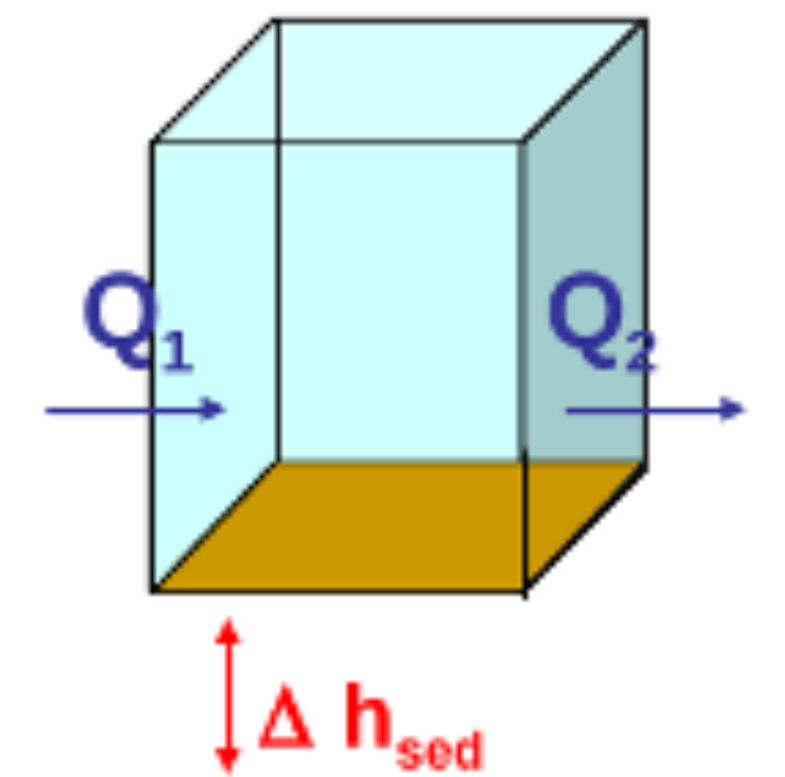
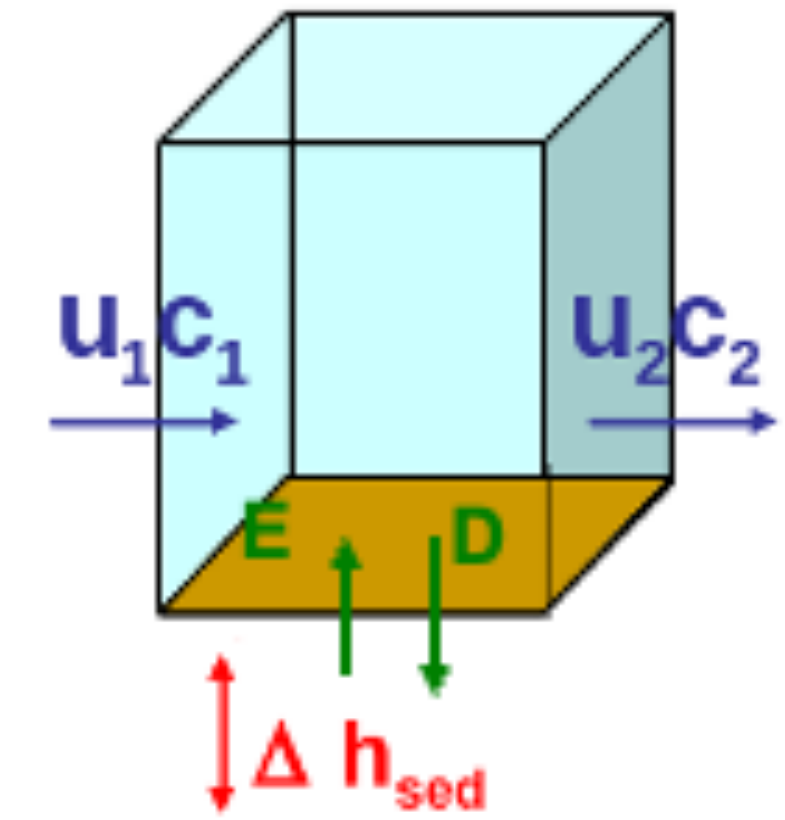
$$\frac{\partial hC}{\partial t} + \frac{\partial hUC}{\partial x} = E - D$$

Bottom evolution: $\frac{\partial h_{sed}}{\partial t} = \frac{1}{C_{sed}} (D - E)$

Models in CROCO

Sediment modeling : main processes in CROCO

- Transport in the water column
- Erosion / deposition
- Bedload transport
- Bed evolution (sand, mud, mixed)
- Morphological evolution



Models in CROCO

Sediment modeling

2 models available :

- USGS model : cpp key SEDIMENT
 - « legacy » model
 - originally included in ROMS-AGRIF
 - available in ROMS-RUTGERS and OAWST

- IFREMER model : cpp key MUSTANG
 - french model
 - originally included in MARS3D
 - available since 1.2 (released Jan 2022)

Conclusion

- Wide range of complex phenomena at different scales
- Drivers : currents, waves and depth
- Complexity => hard to observe and model

- Very fast review, non-exhaustive

Let's dive into CROCO !!

